

The background of the slide is a photograph of ocean waves, showing a mix of deep blue and lighter turquoise water with white foam on the crests. The waves are moving towards the viewer, creating a sense of depth and movement.

The RISE and fall of the Sea

(as seen in satellite altimeter data)

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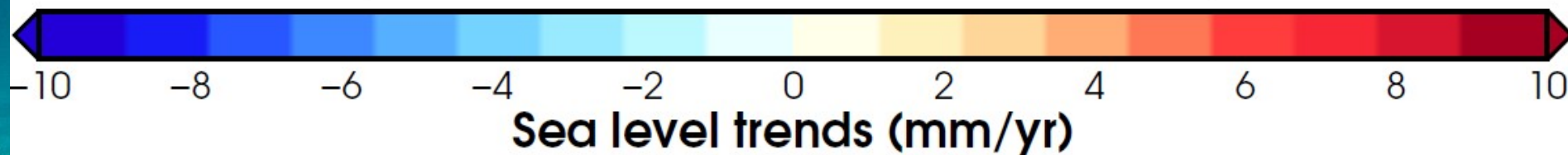
Remko Scharroo (EUMETSAT)

Sok Kuh Kang (KORDI)

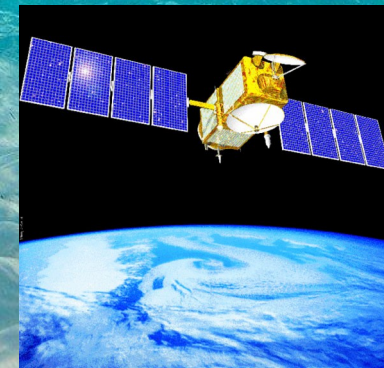
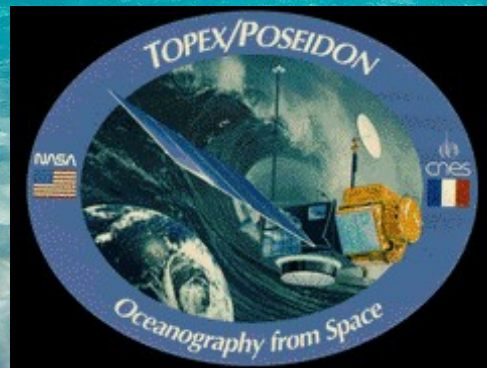
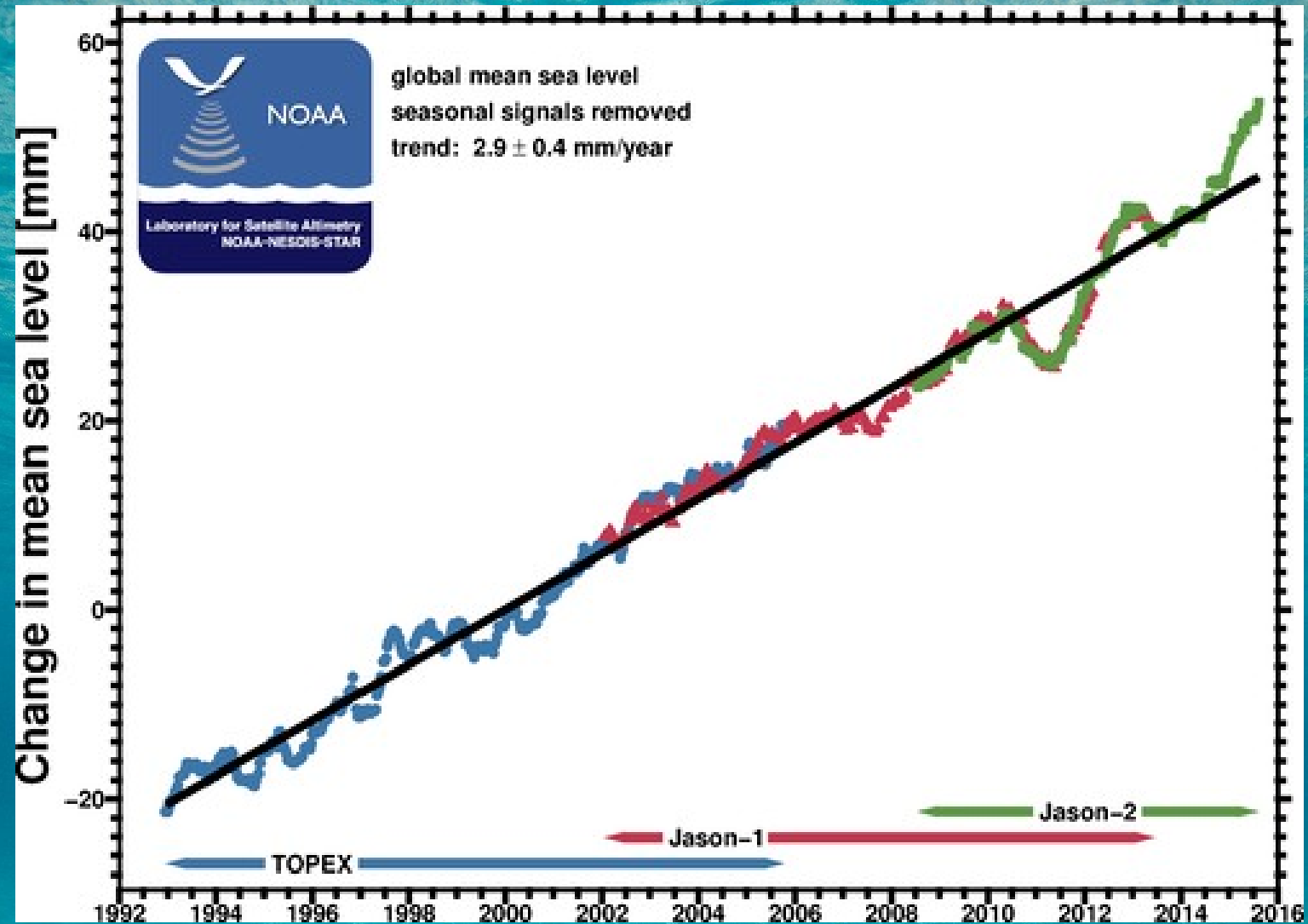
Rob Bell (NIWA)

TOPEX, J1, and J2
1992.96 – 2015.60

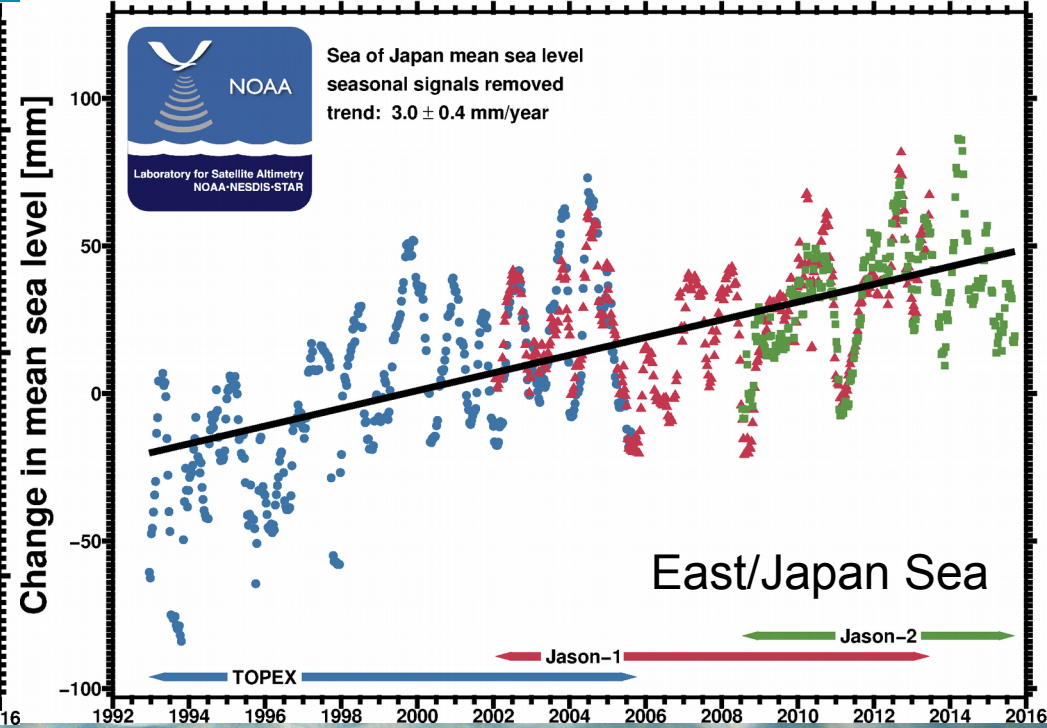
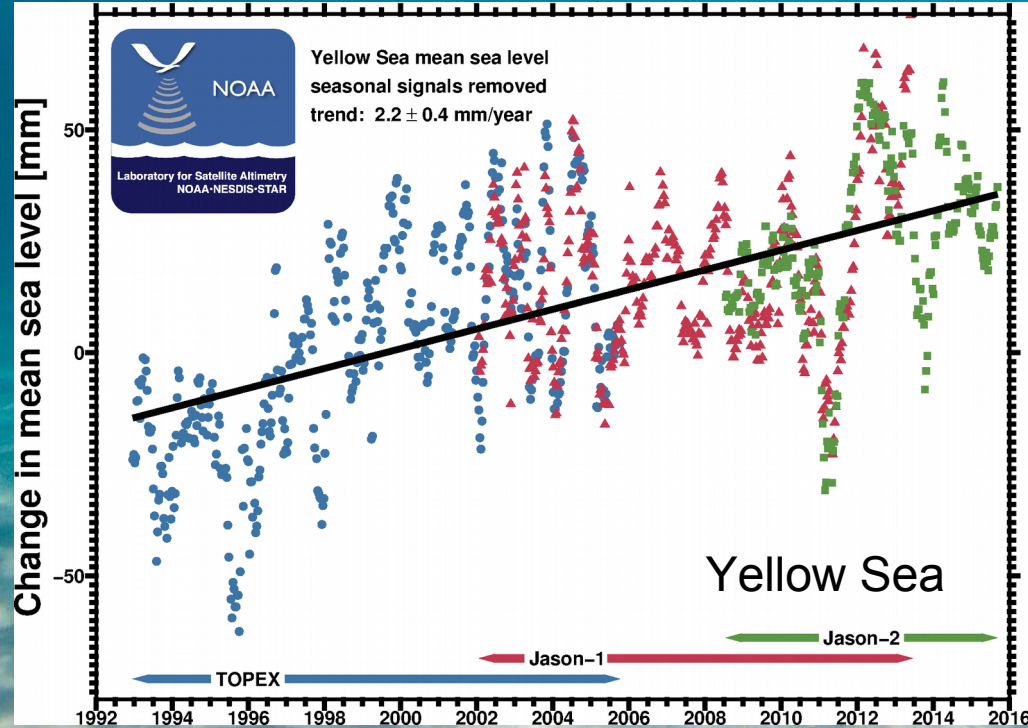
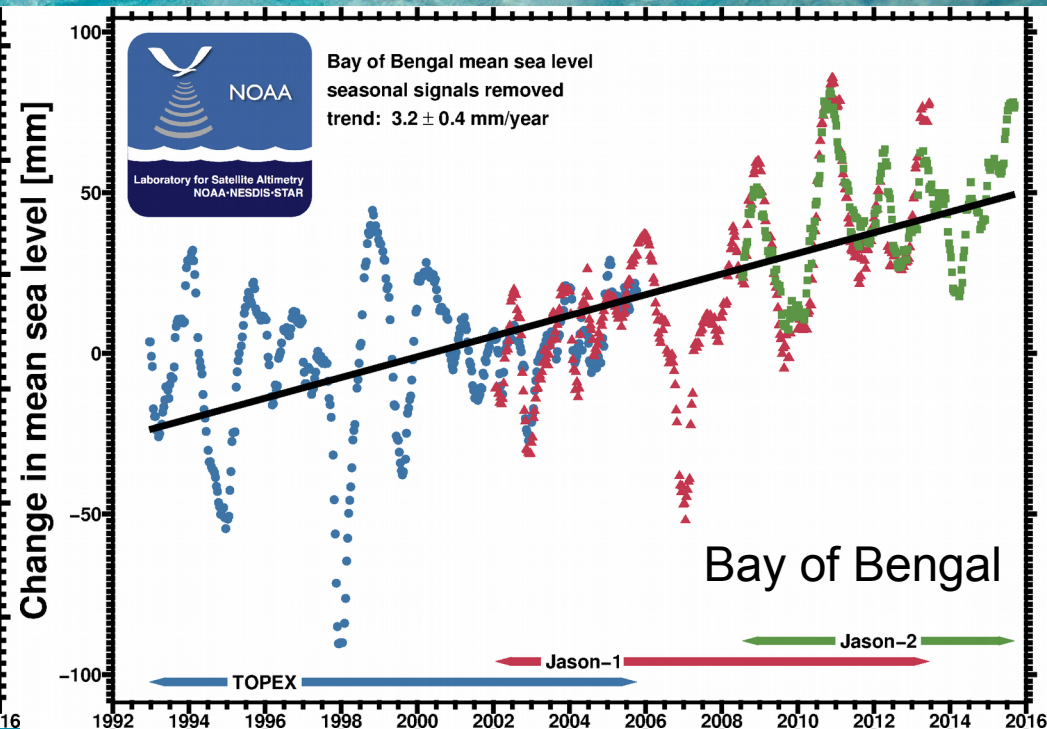
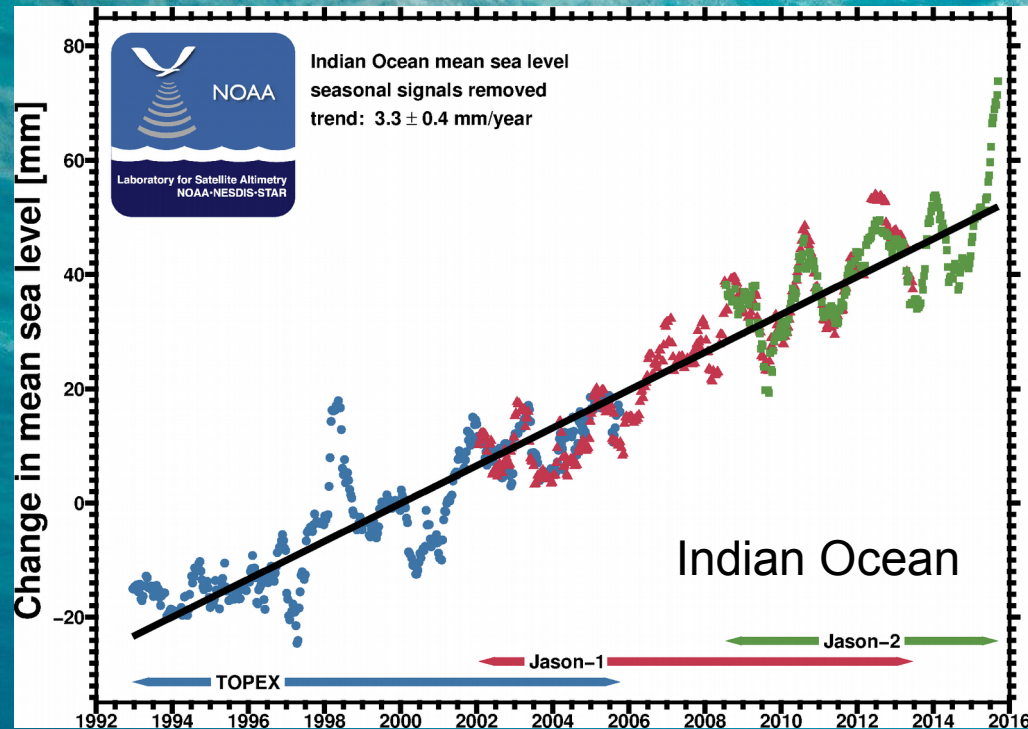
NOAA/Laboratory for Satellite Altimetry



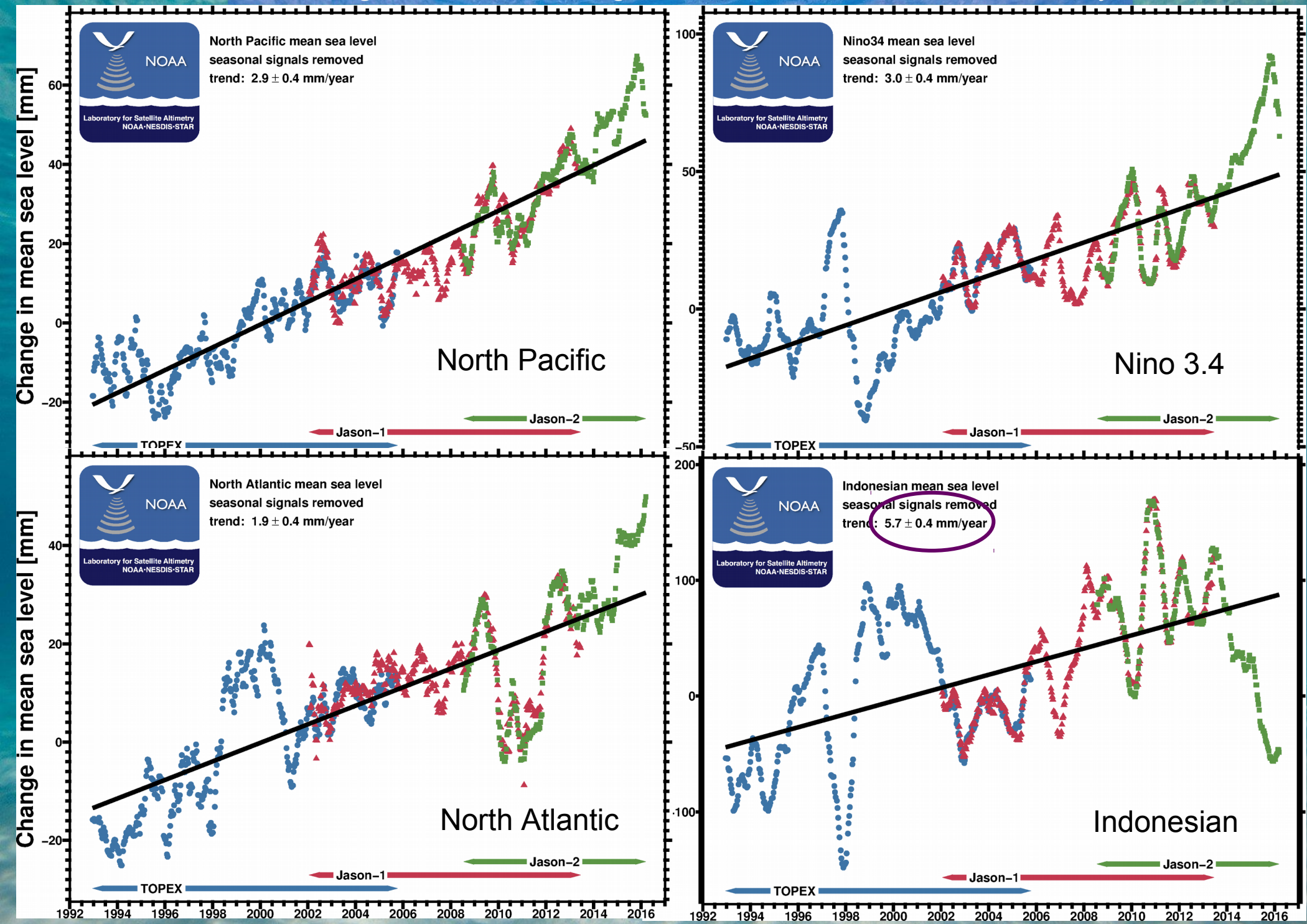
Global mean sea level rise (from NOAA, no GIA applied)



Select regional trends: stronger interannual and decadal variability



Select regional trends: stronger interannual and decadal variability



The aims of this work are to compute local and regional *sea level trends* from TOPEX/Poseidon, Jason-1 and Jason-2 satellite altimeter data while taking into account the important *modes* of climate variability

We include in the analyzes the following modes/indices:

1. Lunar nodal tide (Mn: period = 18.6 years)
2. Low-pass filtered Pacific Decadal Oscillation (PDO) Index
3. High-pass filtered Multivariate ENSO Index (MEI)

The IOS Versatile Analysis (IVA)

we modify the **Versatile Harmonic Analysis** (Foreman et al. 2009)

$$h(t_j) = Z_0 + a(t_j - t_c) + \sum_{k=1}^n f_k(t_j) A_k \cos[V_k(t_j) + u_k(t_j) - g_k] + R(t_j)$$

Trend

to also include non-harmonic climate indices

$$h(t_j) = Z_0 + a(t_j - t_c) + \sum_{i=1}^{N_c} b_i I_i(t_j) + \sum_{k=1}^n f_k(t_j) A_k \cos[V_k(t_j) + u_k(t_j) - g_k] + R(t_j)$$

N_c = number of climate indices (I_i) and n = number of harmonic terms

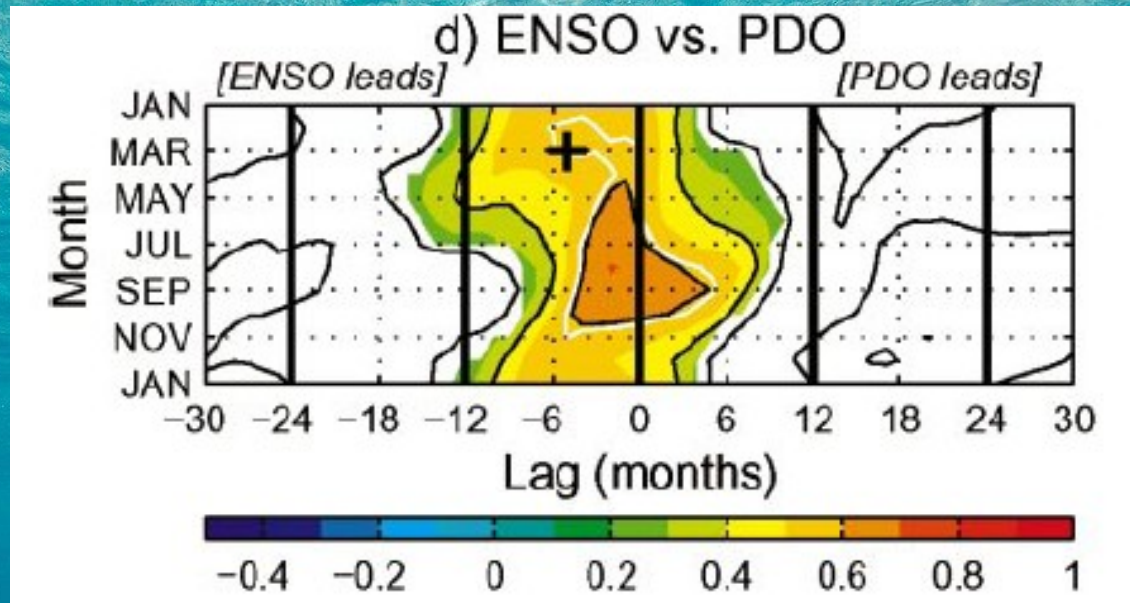
We use SVD to solve for Z_0 , a , b_i ($i=1, N_c$), A_k and g_k ($k=1, n$)

with corresponding constituent error covariances and correlations:

$$r_{ij} = C_{ij} / \sqrt{C_{ii} C_{jj}}$$

(Cherniawsky et al. 2001, 2010; Foreman et al. 2009)

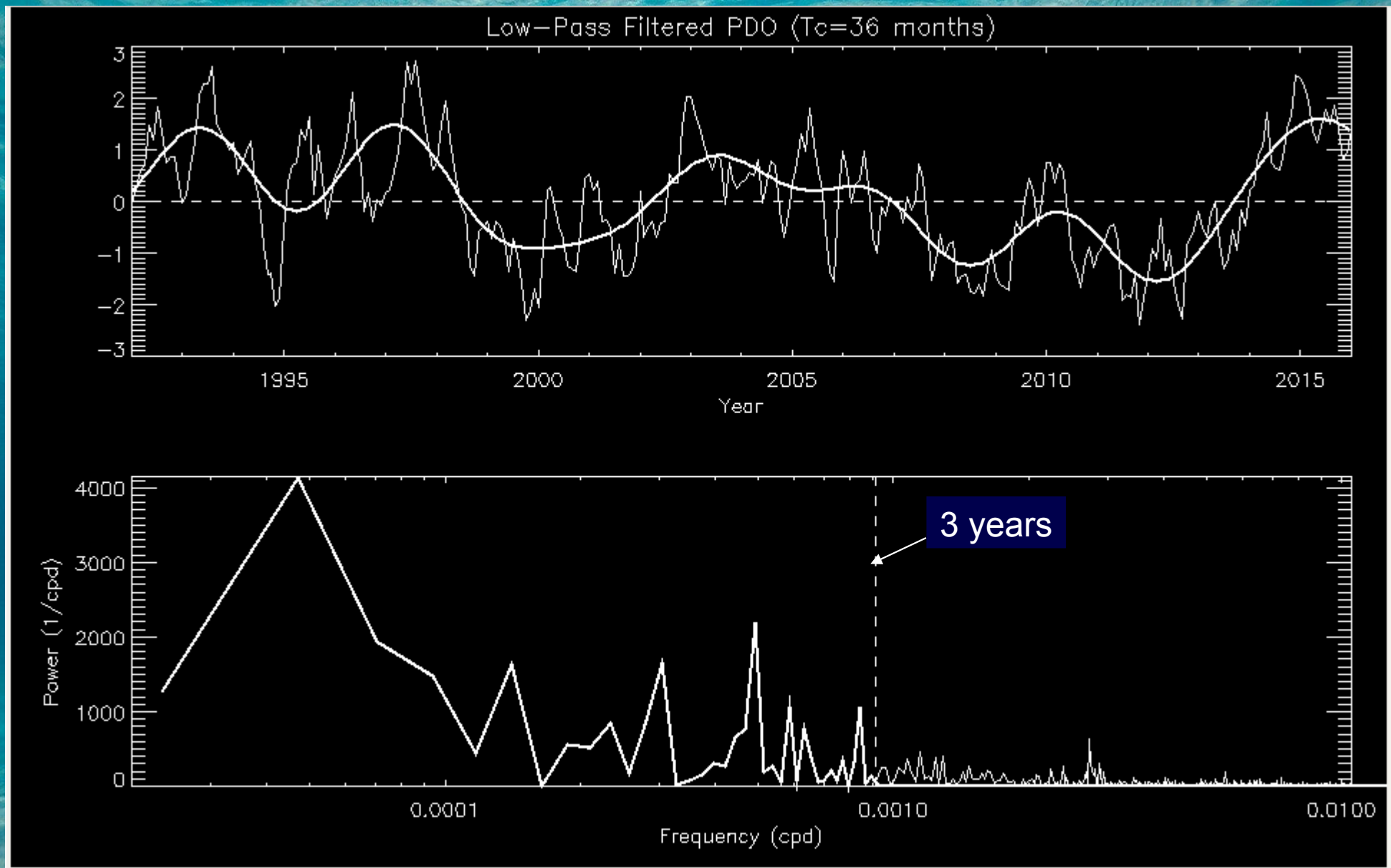
But ENSO (MEI) and PDO are not independent!



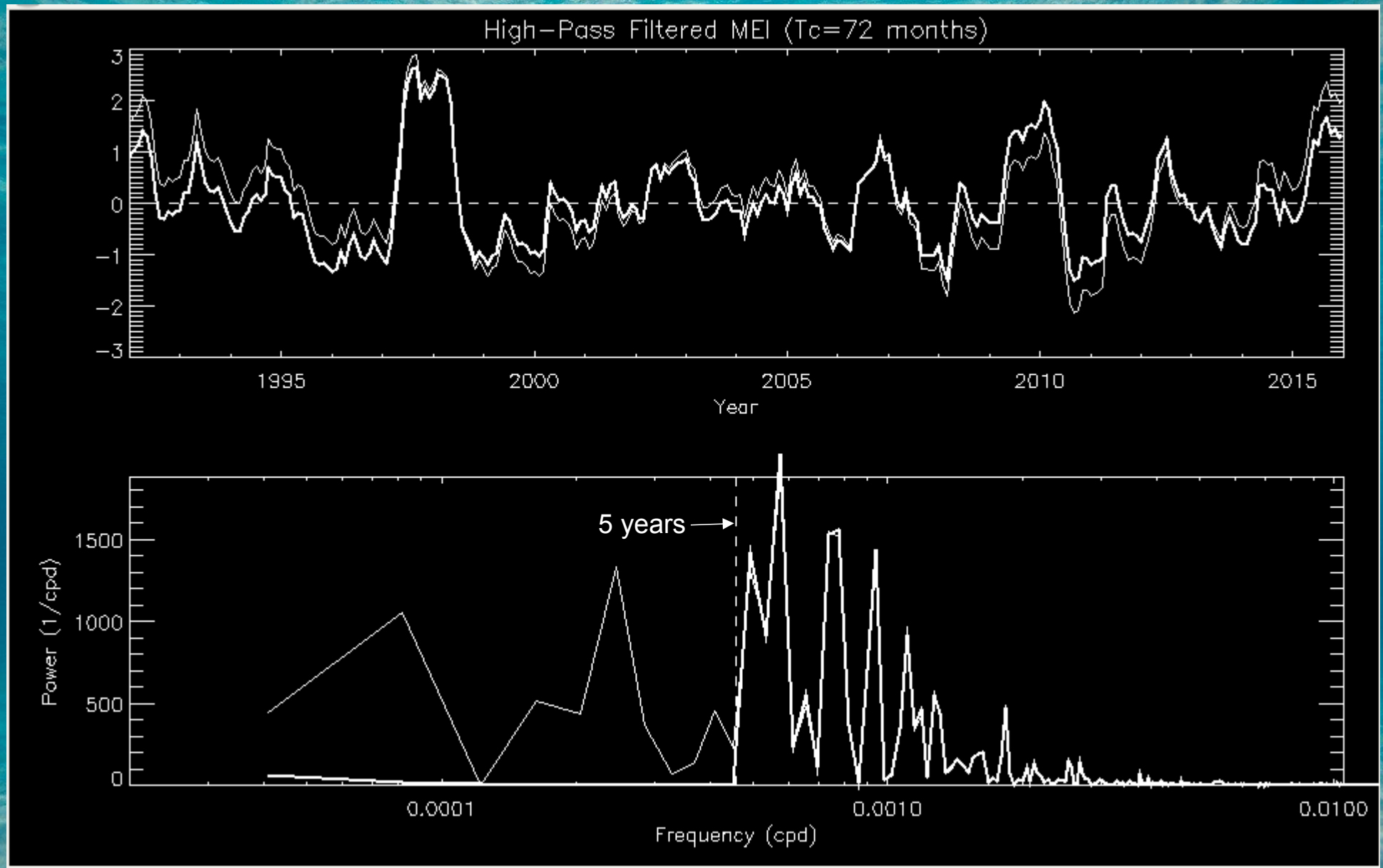
(from Newman et al., *J. Climate* 2003)

Therefore, when we use MEI and PDO indices 'as is'
their analysed amplitudes are correlated, with $r \sim 0.5$!
We therefore choose to filter these indices.

Input: Unfiltered & Low-Pass Filtered PDO Index



Input: Unfiltered & High-Pass Filtered MEI



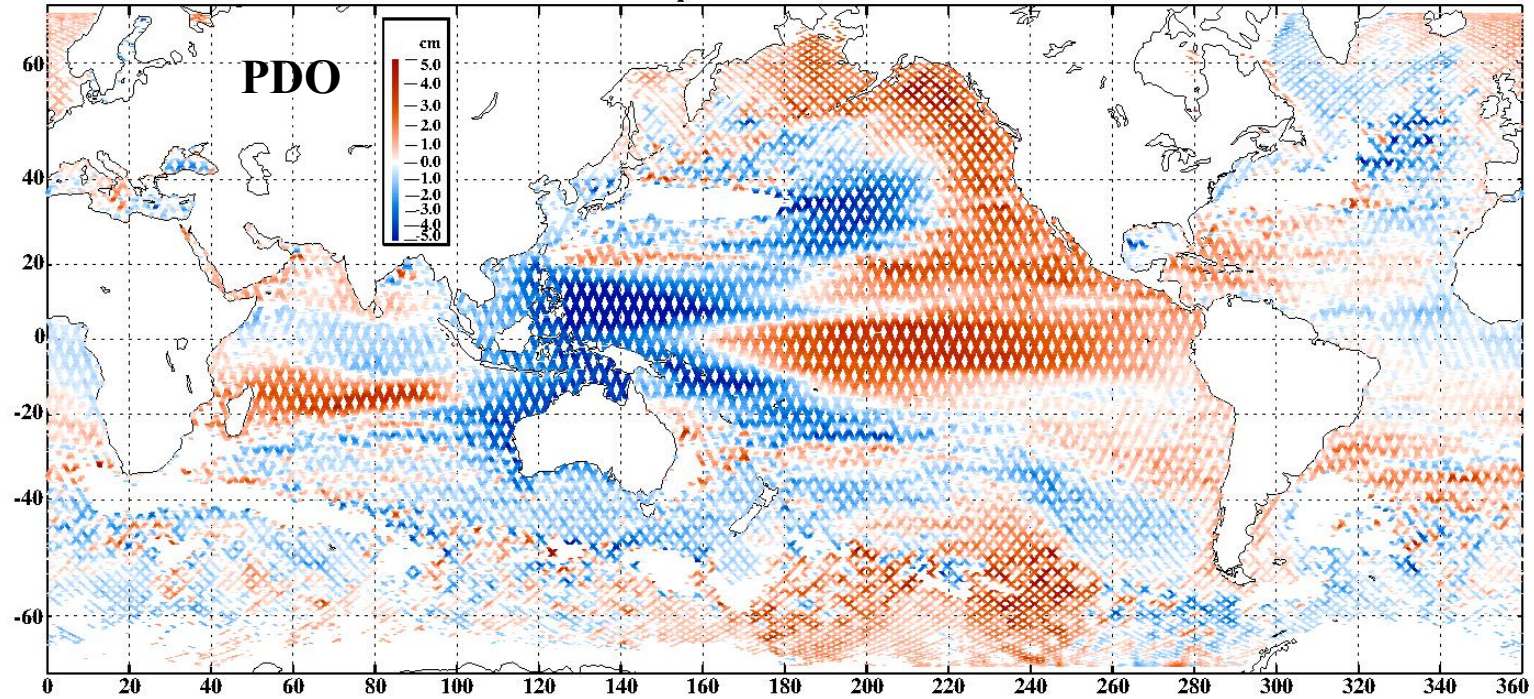
Trend is computed

```
31 1 1
Z0      0.0000000000
PDO     0.0000000000  PDO index
MEI     0.0000000000  MEI index
Mn      0.0000061290  18.6-yr
SA      0.0001140741
SSA     0.0002281591
MM      0.0015121518
MF      0.0030500918
Q1n     0.0372123736  nodal sat
Q1      0.0372185026
RHO1    0.0374208736
O1n     0.0387245256  nodal sat
O1      0.0387306544
NO1     0.0402685944
P1      0.0415525871
S1      0.0416666721
K1      0.0417807462
K1n     0.0417868751  nodal sat
PSI1    0.0418948203
J1      0.0432928981
2N2     0.0774870970
MU2     0.0776894680
N2      0.0789992488
NU2     0.0792016198
M2n     0.0805052717  nodal sat
M2      0.0805114007
L2      0.0820235525
T2      0.0832192592
S2      0.0833333333
K2      0.0835614924
K2n     0.0835676239  nodal sat
23      09      92      28      01      16      19      20
```

An example of the
IVA list of constituents
(*tide14_rv.dat*)

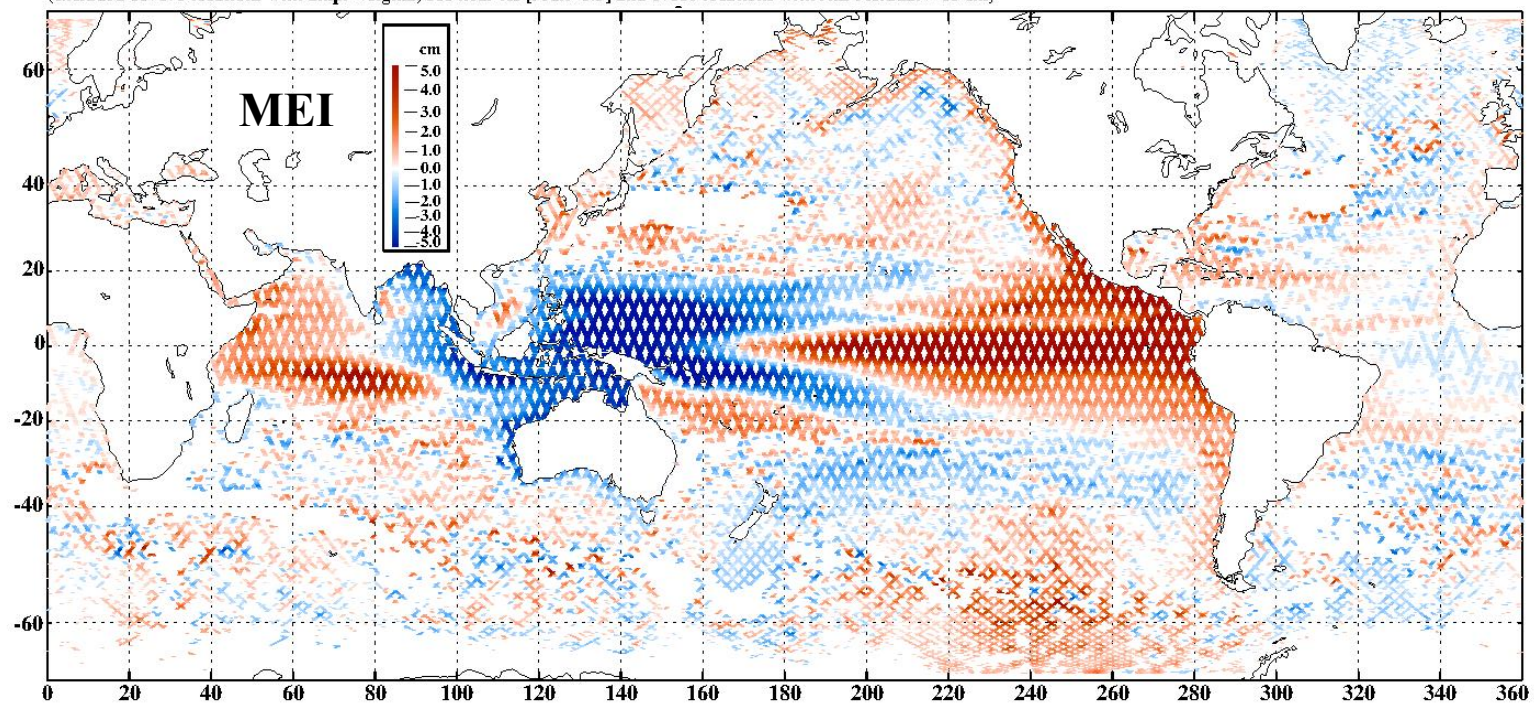
Analysis results at a single TPJ location off Vancouver Island (48.89N, 126.01W)

		AMPLITUDE	PHASE LAG	Amp SD Est.	t-test value
Z0	0.000000000	-0.01180	0.000	0.00383	3.08
Trend	0.000000000	0.00168	0.000	0.00159	1.06
PDO	0.000000000	0.02194	0.000	0.00486	4.51
MEI	0.000000000	0.01881	0.000	0.00701	2.68
Mn	0.000006129	0.00876	342.579	0.00732	1.20
SA	0.000114074	0.08616	2.380	0.00741	11.63
SSA	0.000228159	0.02210	237.800	0.00668	3.31
MM	0.001512152	0.00587	171.076	0.00616	0.95
MF	0.003050092	0.01513	170.605	0.00622	2.43
Q1n	0.037212372	0.00540	279.916	0.00563	0.81
Q1	0.037218504	0.04454	216.482	0.00689	6.47
RHO1	0.037420873	0.00227	261.364	0.00612	0.37
Q1n	0.038724527	0.05064	226.418	0.00624	8.11
Q1	0.038730655	0.25623	226.516	0.00670	38.24
NO1	0.040268596	0.01620	240.833	0.00576	2.81
P1	0.041552588	0.12720	235.380	0.00679	18.73
S1	0.041666672	0.01374	244.933	0.00904	1.52
K1	0.041780747	0.42079	240.437	0.00675	62.30
K1n	0.041786876	0.04631	248.442	0.00714	6.49
PSI1	0.041894820	0.00269	313.953	0.00661	0.41
J1	0.043292899	0.01961	233.485	0.00628	3.12
2N2	0.077487096	0.02004	186.094	0.00630	3.18
MU2	0.077689469	0.02215	181.148	0.00623	3.56
N2	0.078999251	0.19707	210.660	0.00650	30.30
NU2	0.079201616	0.04013	209.014	0.00632	6.35
M2n	0.080505274	0.04017	232.857	0.00638	6.29
M2	0.080511399	0.95505	236.550	0.00679	140.57
L2	0.082023554	0.02026	248.364	0.00631	3.21
T2	0.083219260	0.01922	247.112	0.00638	3.01
S2	0.083333336	0.27314	264.767	0.00638	42.79
K2	0.083561495	0.06720	255.493	0.00680	9.89
K2n	0.083567627	0.02202	250.436	0.00690	3.19
1 largest correlation coef. =	0.2936	at (i,j)=	6	2 for Mn	and Trend
2 largest correlation coef. =	0.2834	at (i,j)=	4	3 for MEI	and PDO
3 largest correlation coef. =	0.2427	at (i,j)=	5	3 for Mn	and PDO
4 largest correlation coef. =	0.2034	at (i,j)=	3	2 for PDO	and Trend
5 largest correlation coef. =	0.1847	at (i,j)=	7	4 for SA	and MEI
6 largest correlation coef. =	0.1729	at (i,j)=	6	5 for Mn	and Mn
7 largest correlation coef. =	0.1506	at (i,j)=	5	2 for Mn	and Trend
8 largest correlation coef. =	0.1404	at (i,j)=	18	15 for Q1	and Q1n



Min, Max, Mean, StDev: -10.56, 10.09, 0.05, 1.95 cm (N=351534) Mean sig = 0.45 cm
(excluded 187175 locations with ampl < sigma, 111 near Xs [redit=0.5] and 17150 locations with rms residual > 15 cm)

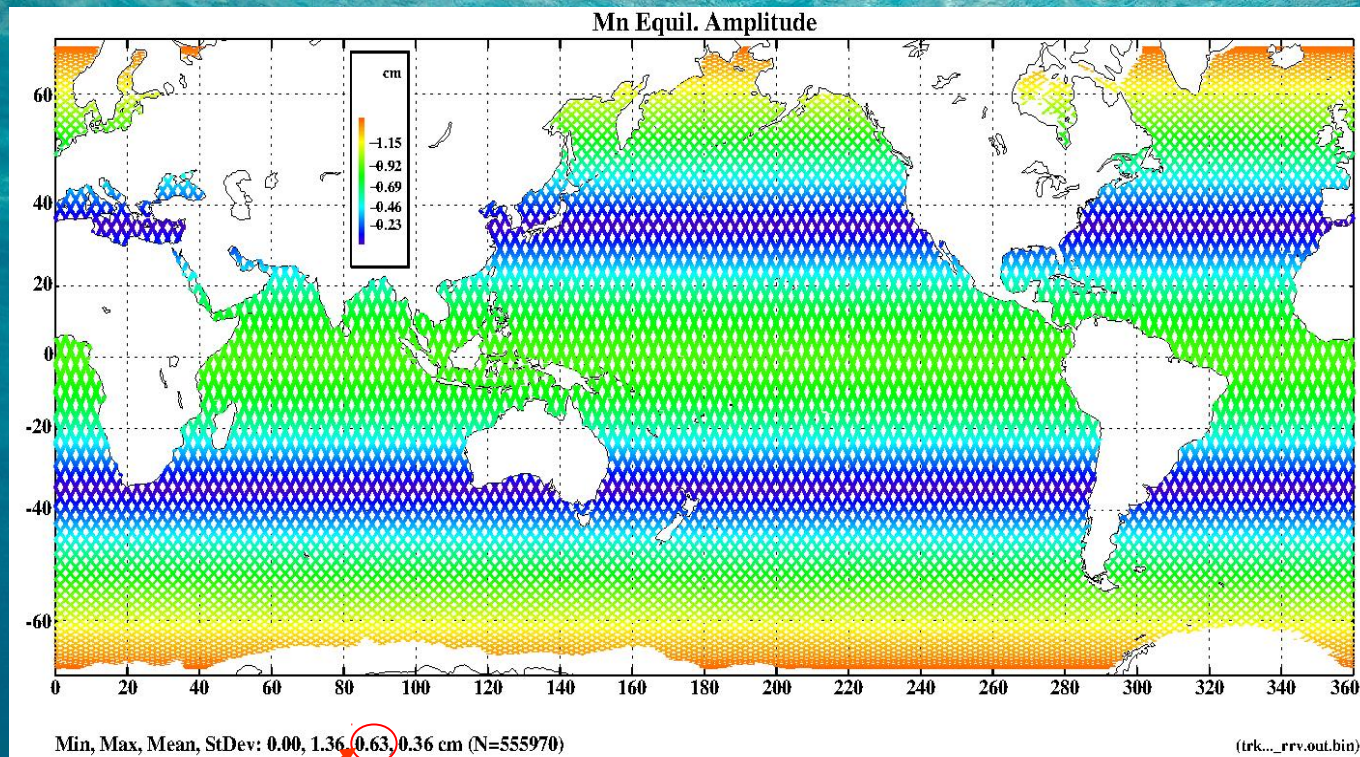
(trk...rrv.out.bin)



Min, Max, Mean, StDev: -9.54, 8.08, 0.33, 2.31 cm (N=258275) Mean sig = 0.62 cm
(excluded 287439 locations with ampl < sigma and 10256 locations with rms residual > 15 cm)

(trk...rrv.out.bin)

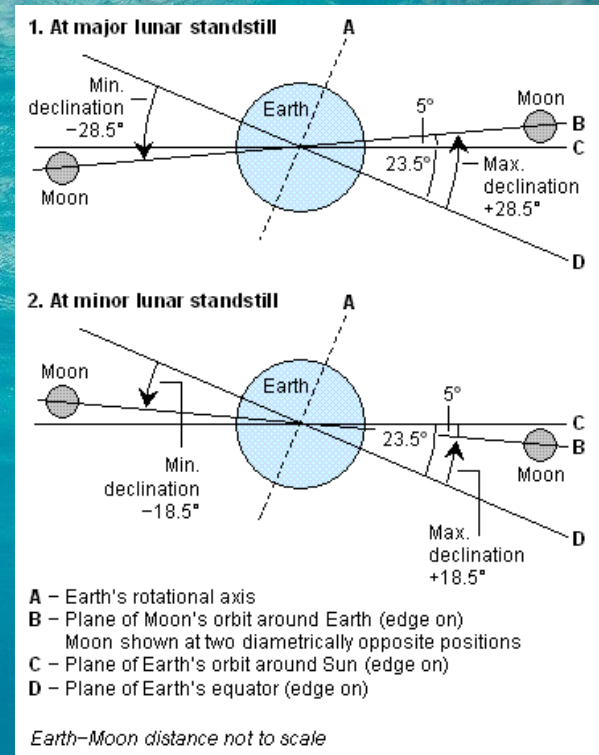
Theoretical Amplitude of the 18.6-year Lunar Nodal Tide

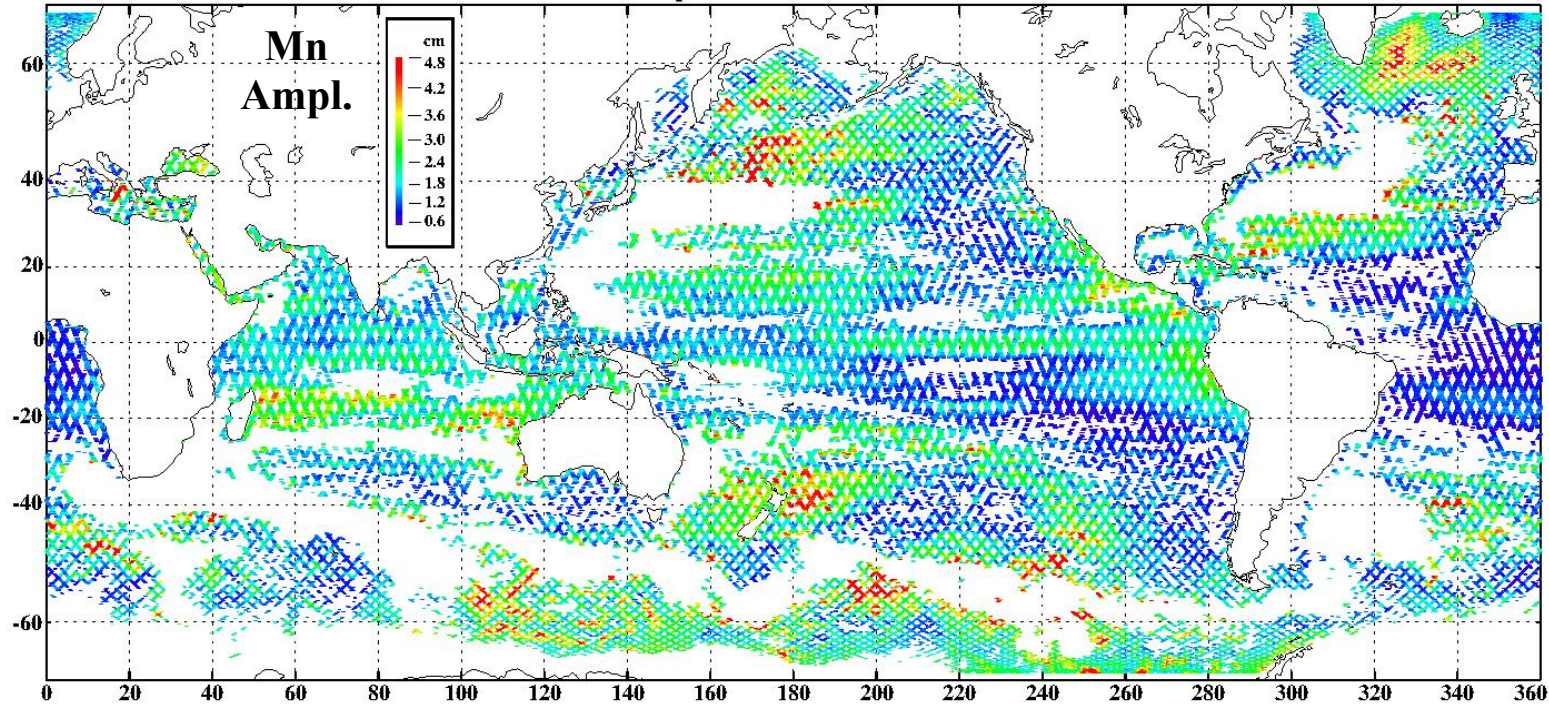


Mean = 0.63 cm

$$H_N = 18 \left(\frac{3}{2} \sin^2 \varphi - \frac{1}{2} \right) \cos N \quad (\text{in mm})$$

N is the mean longitude of the Moon's ascending node
Last maximum of $\cos N$ (when $N=0$) was in Nov. 2006



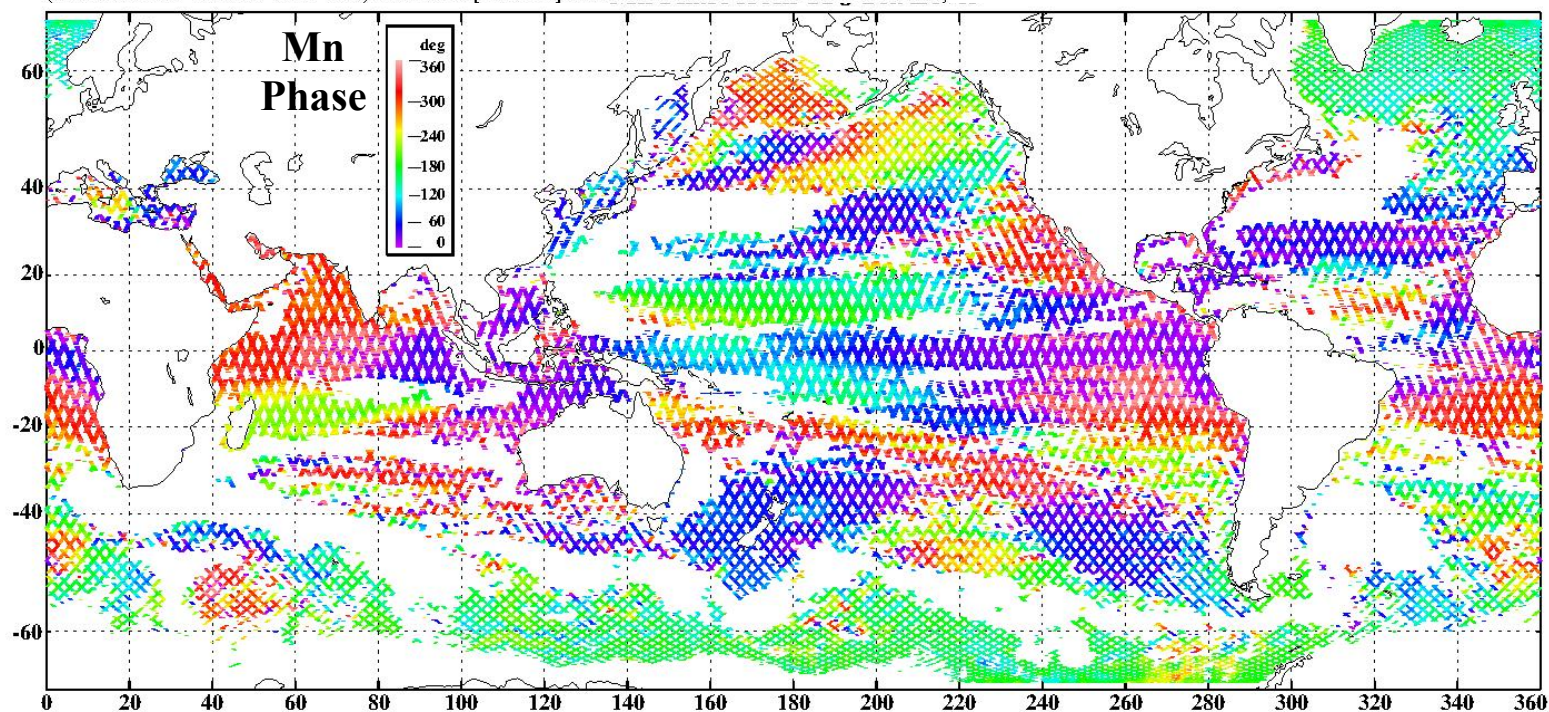


Analyzed
18.6-year
lunar nodal
tide (Mn)

Mean
amplitude
is 2.0 cm,
i.e. ~3 times
larger than
theoretical

Min, Max, Mean, StDev: 0.55, 11.23, 2.00, 0.87 cm (N=291834) Mean sig = 0.56 cm
(excluded 208337 locations with $r < 2.0$, 118 near Xs [redit=0.5] and 55681 locations with rms residual > 10 cm)

(trk..._rrv.out.bin)



Mn phase
& amplitude
have strong
regional
variations

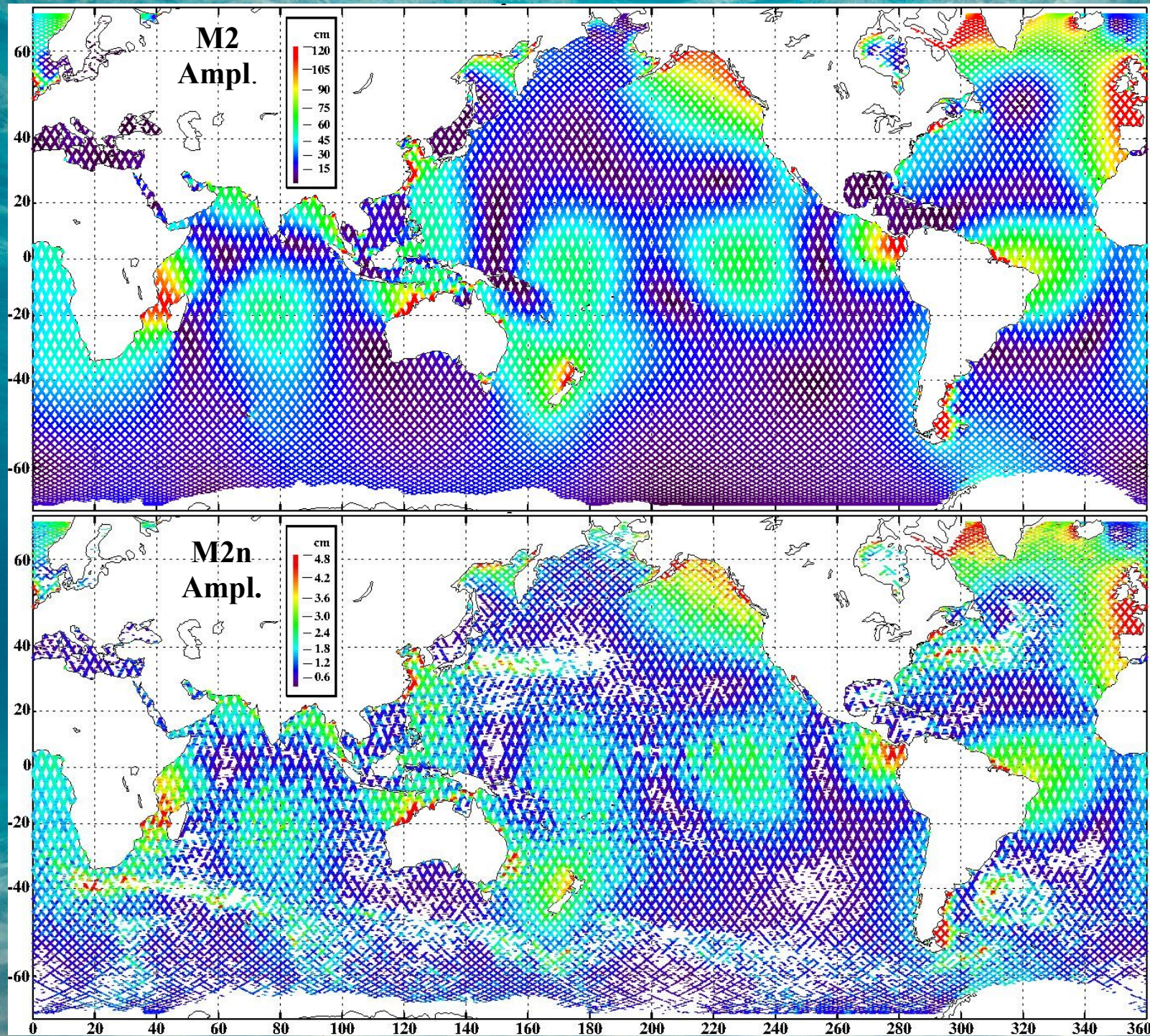
and

do not look
at all like the
theoretical Mn!

Min, Max, Mean, StDev: 0.0, 359.0, 163.8, 110.3 deg (N=291834) Mean sig = 0.56 cm

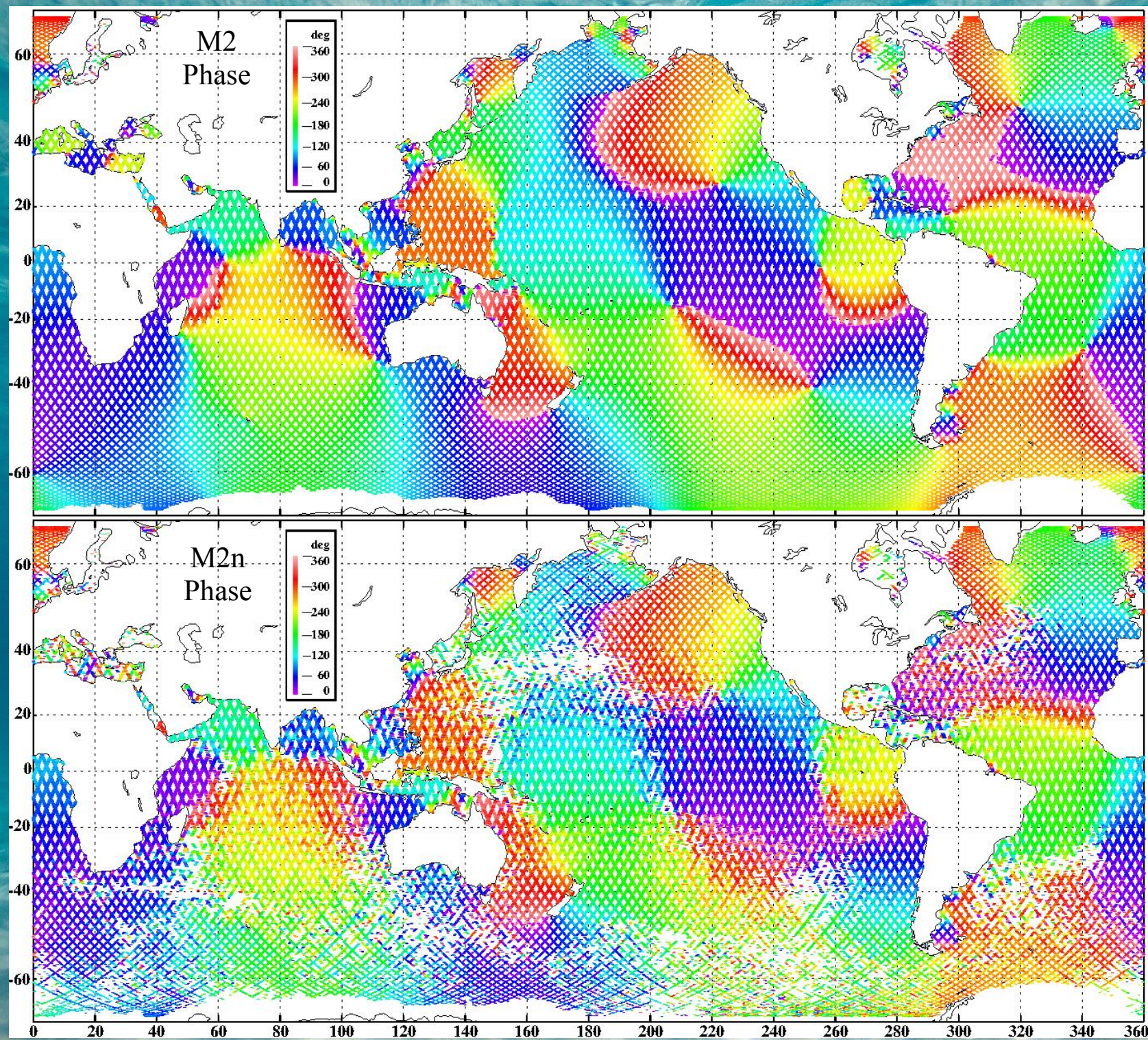
(trk..._rrv.out.bin)

Lunar nodal satellites



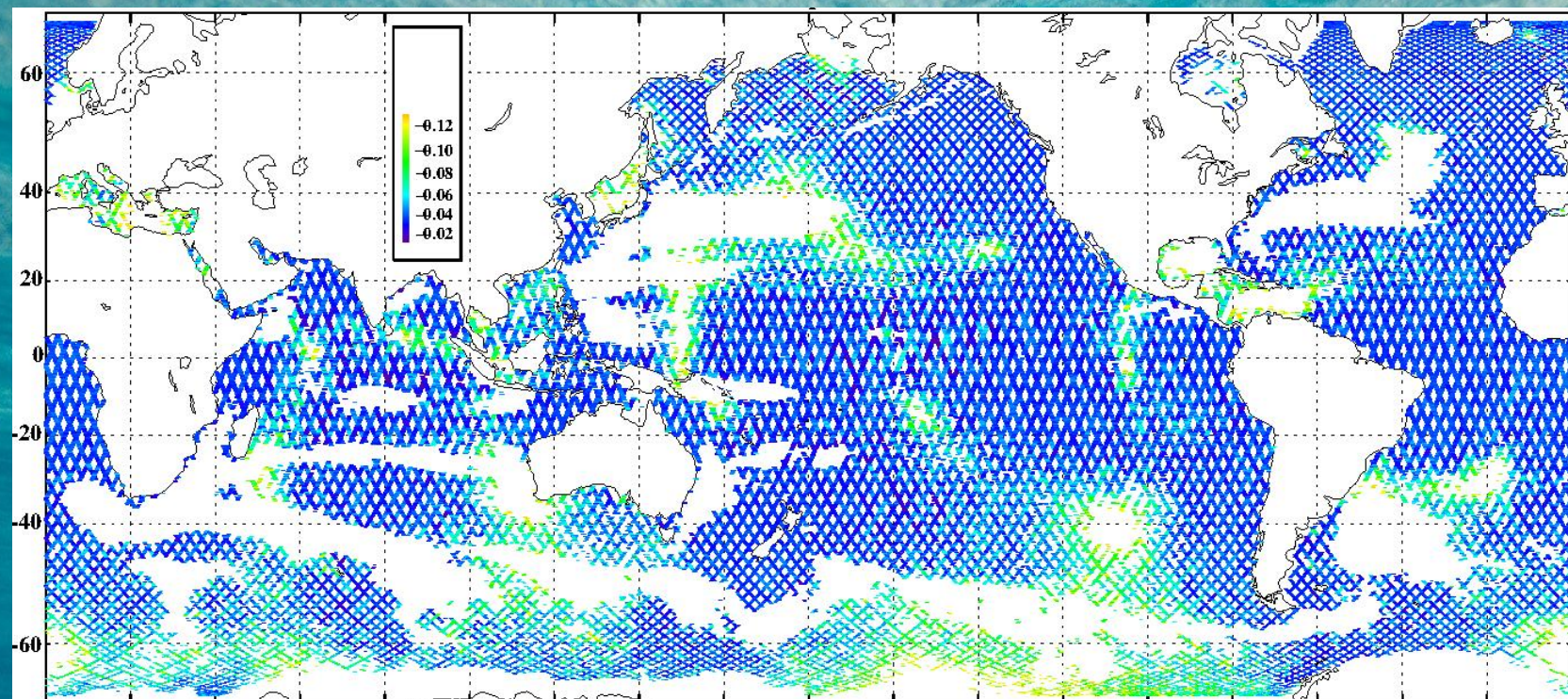
Analyzed
M2 & M2n
amplitudes

Note the
different
amplitude
scales!



Analyzed
M2 & M2n
phases:

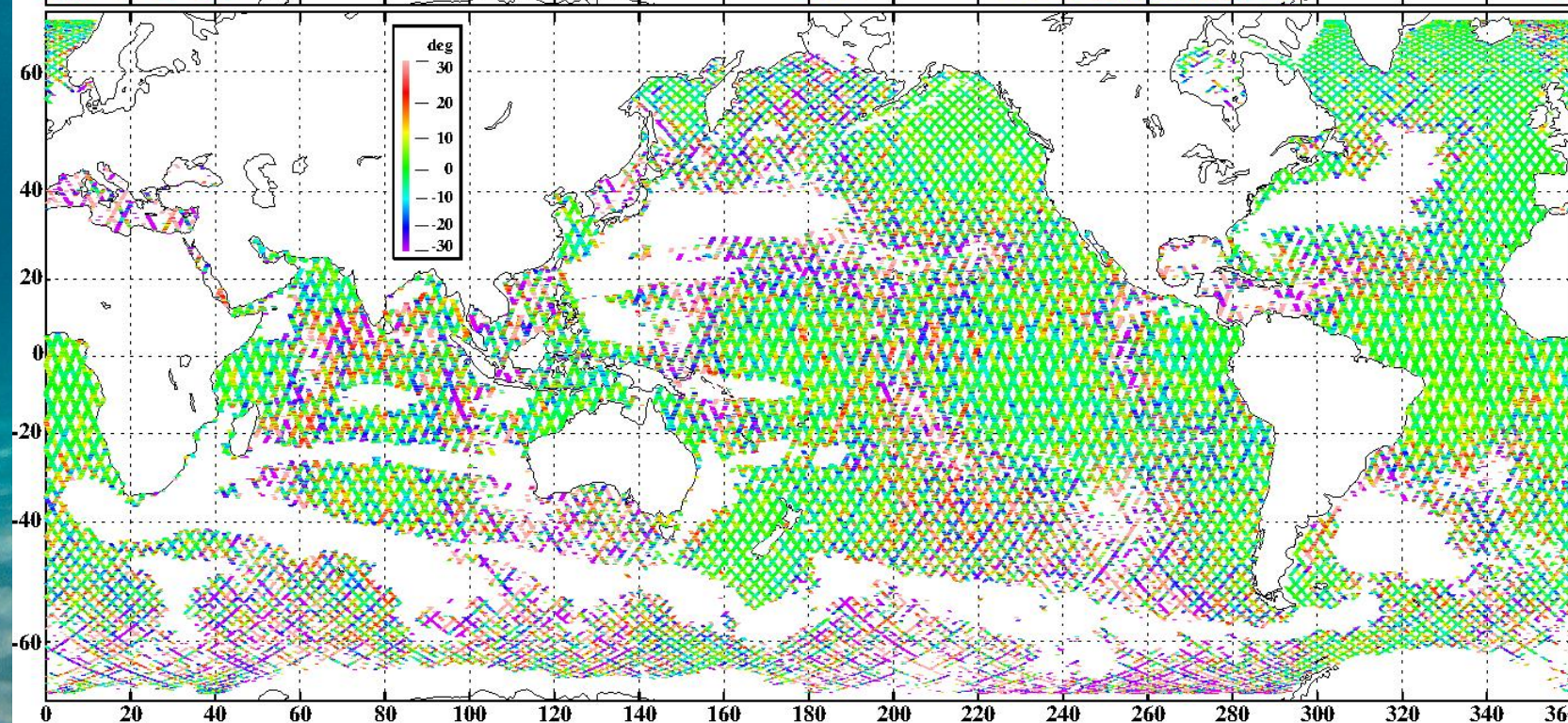
They are
in phase!



M2n/M2
amplitude
ratio

median:
4.0%

theory:
3.7%



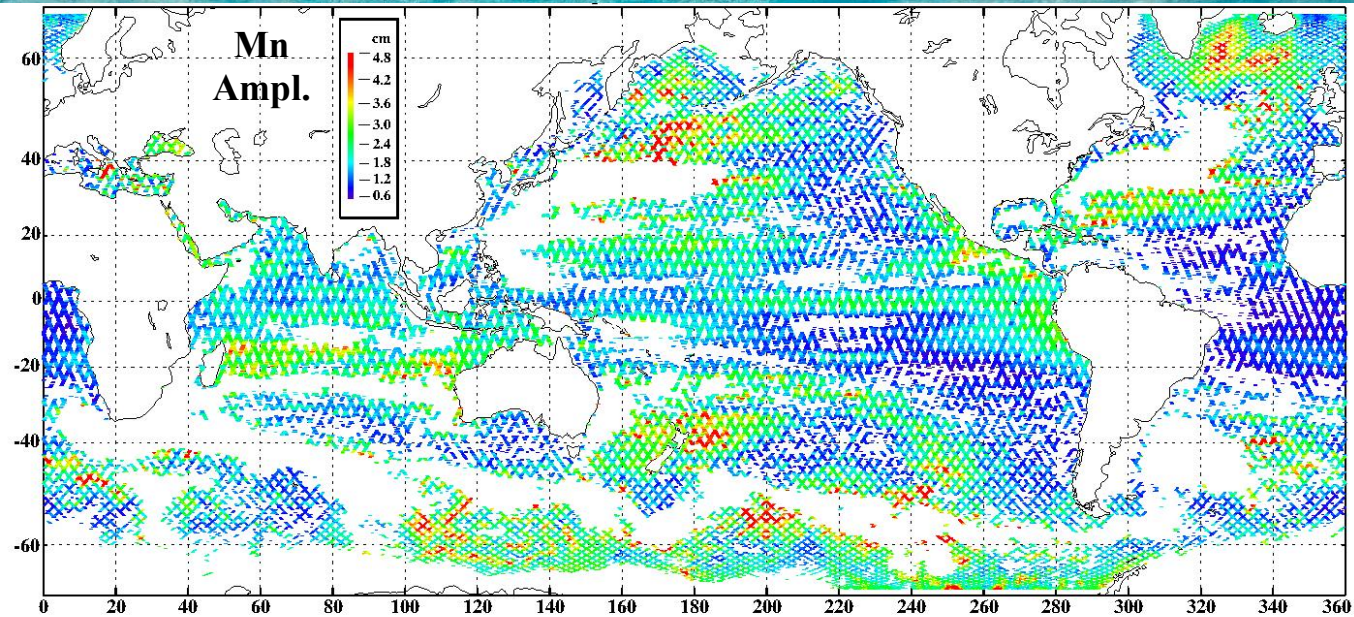
M2n-M2
phase
difference

median:
0.9 deg

theory:
0.0 deg

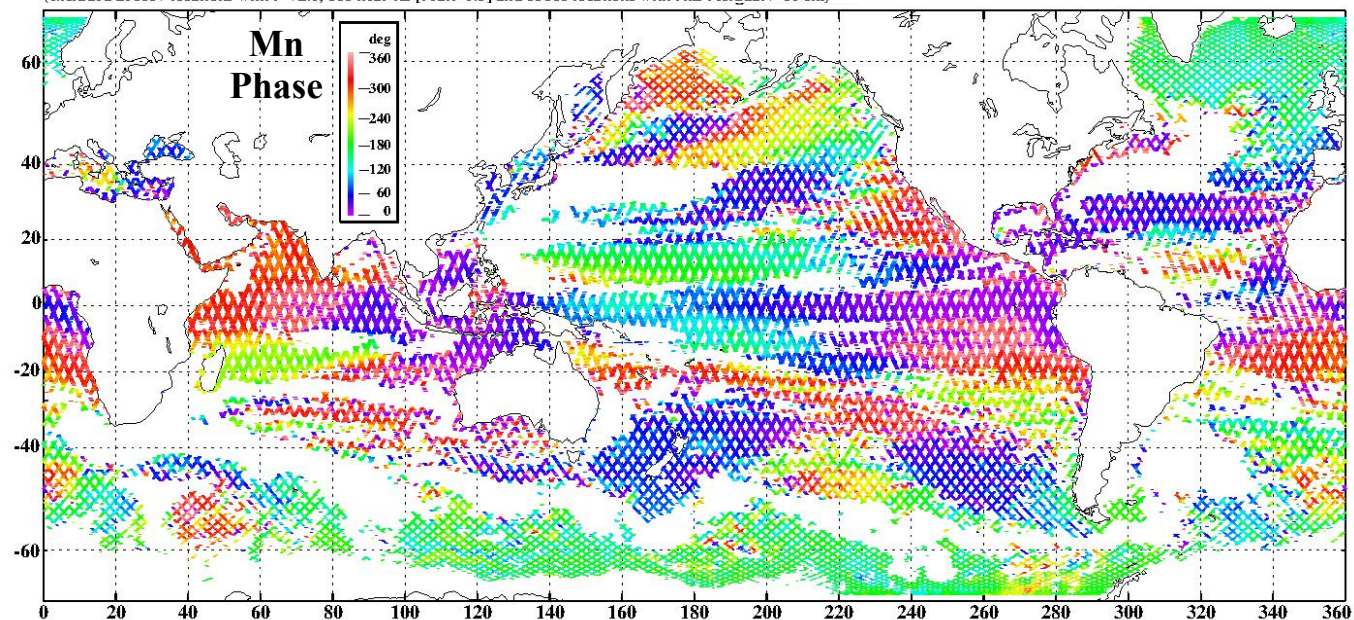
(see also Cherniawsky et al., *Cont. Shelf Res.* 2010)

A conclusion: the 18.6-year 'tide' is generated from modulation of the diurnal and semi-diurnal tides by lunar nodal satellites



Min, Max, Mean, StDev: 0.55, 11.23, 2.00, 0.87 cm (N=291834) Mean sig = 0.56 cm
(excluded 208337 locations with $r < 2.0$, 118 near Xs [redit=0.5] and 55681 locations with rms residual > 10 cm)

(trk..._rrv.out.bin)

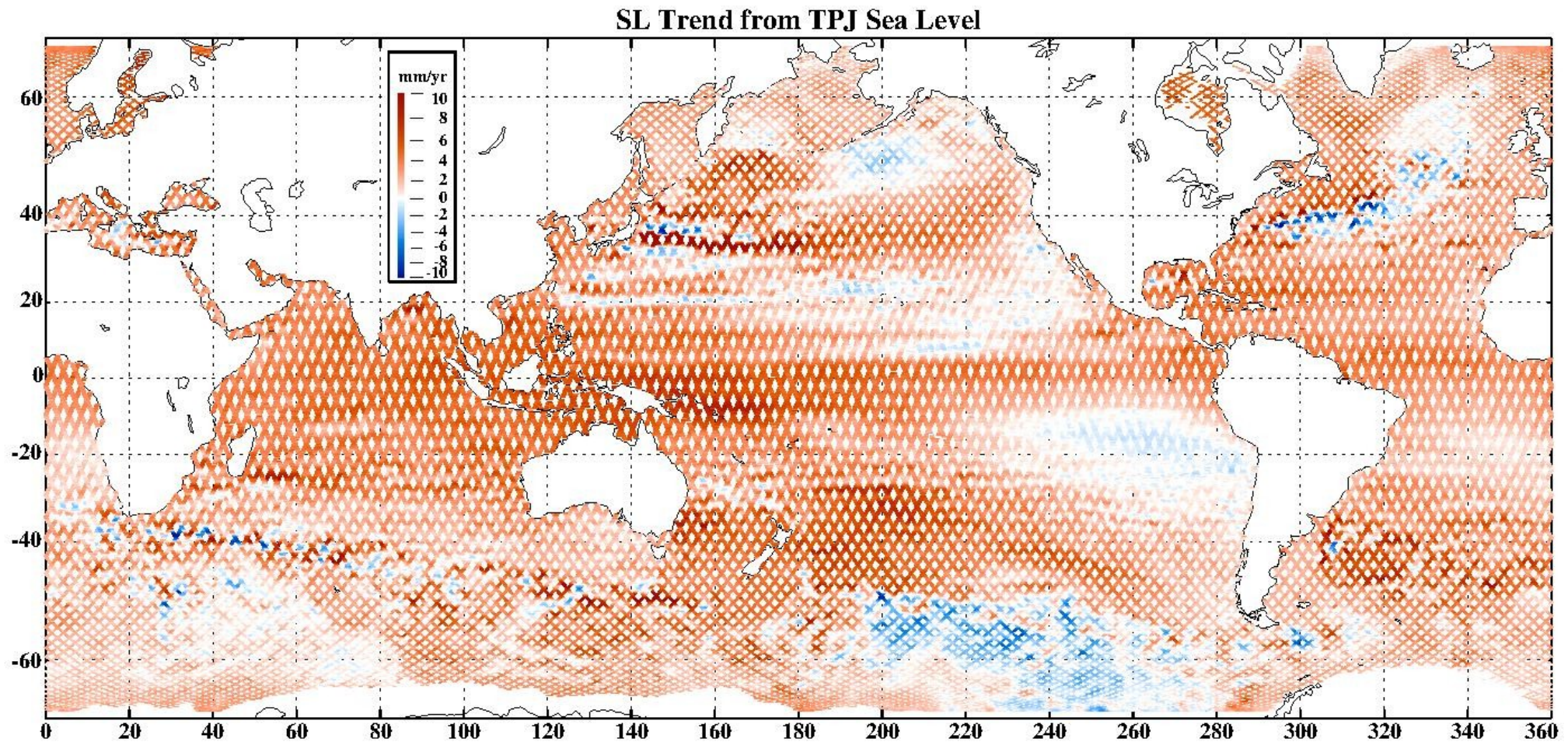


Min, Max, Mean, StDev: 0.0, 359.0, 163.8, 110.3 deg (N=291834) Mean sig = 0.56 cm

(trk..._rrv.out.bin)

(for 18.6-year tide in Bering Sea, see Foreman et al., *J. Mar. Res.* 2006)

Analyzes results: Linear Trend (Sep. 1992 to Jan. 2016)



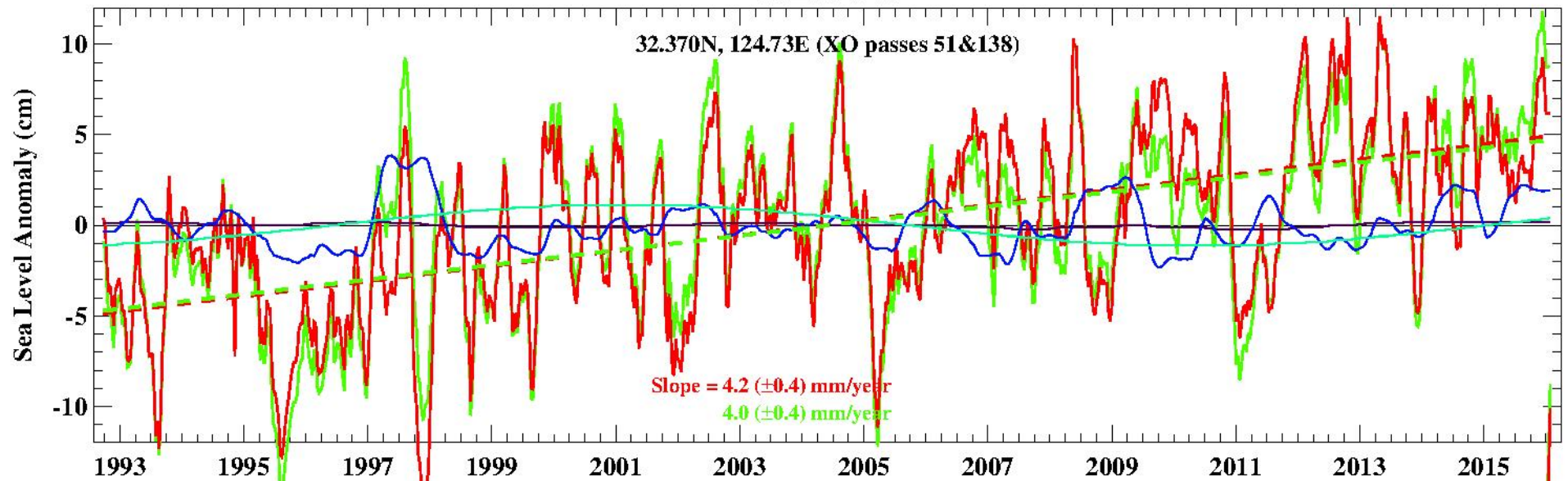
Average trend = 2.8 mm/yr (without GIA)
Its average error = 1.25 mm/yr
(N=555,970 TPJ locations)



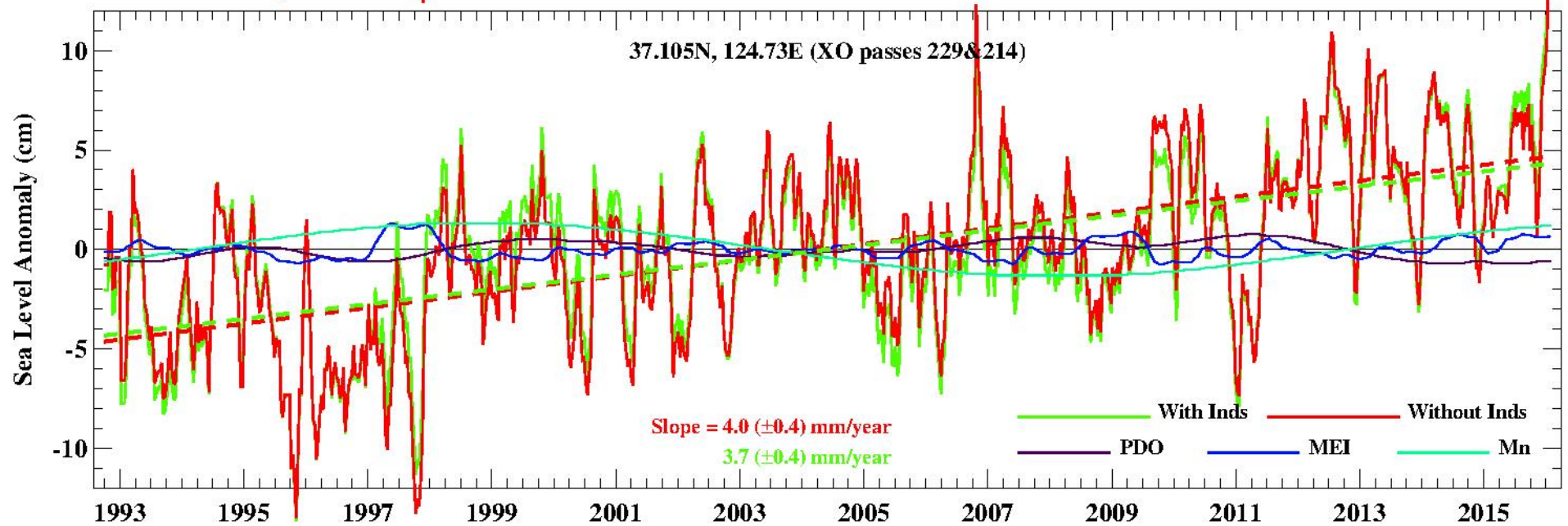
We repeated these analyzes at TPJ pass cross-overs
and inspect the effects of climate indices on
sea level trends at select cross-over locations

Yellow Sea

32.370N, 124.73E (XO passes 51&138)

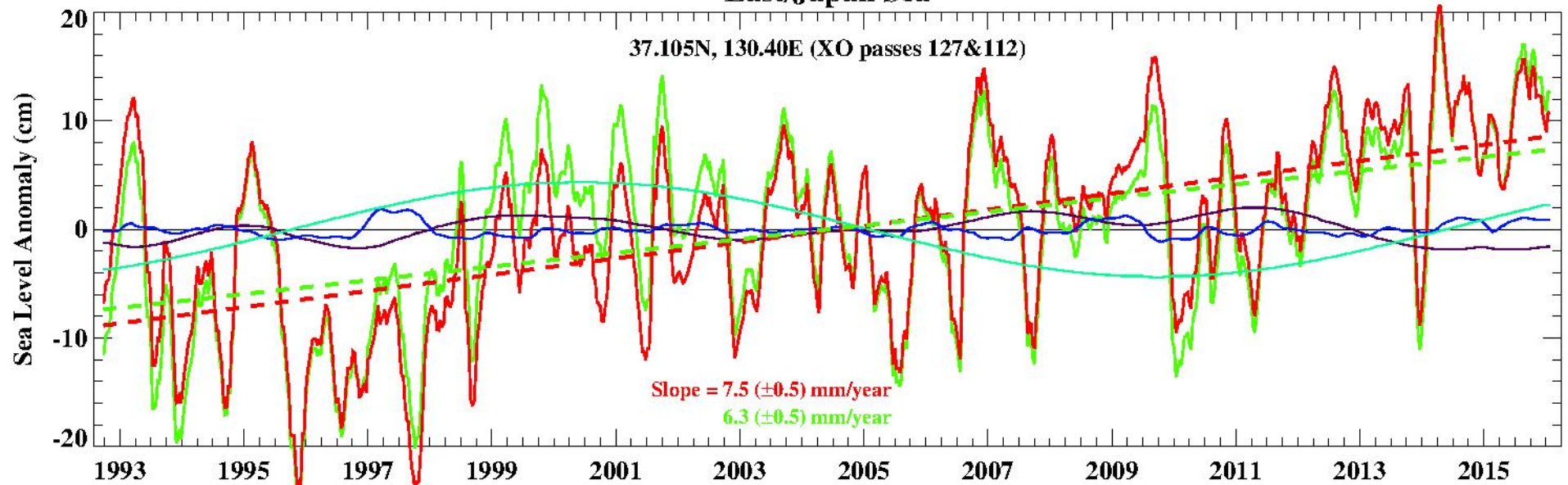


37.105N, 124.73E (XO passes 229&214)

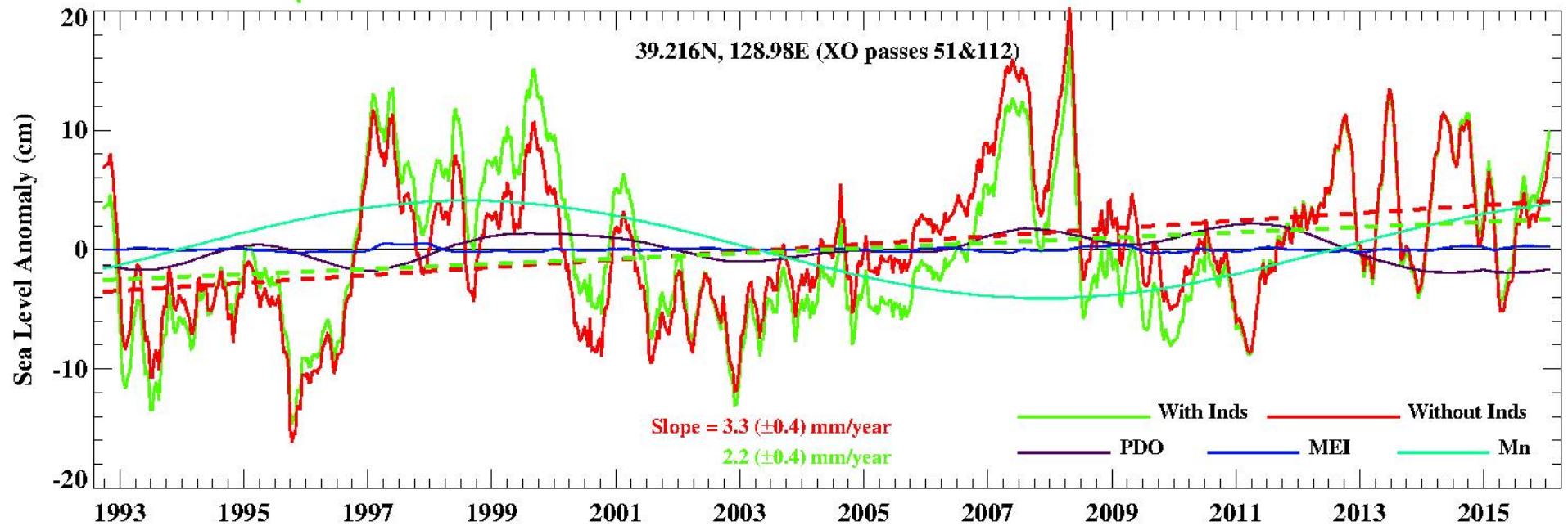


East/Japan Sea

37.105N, 130.40E (XO passes 127&112)

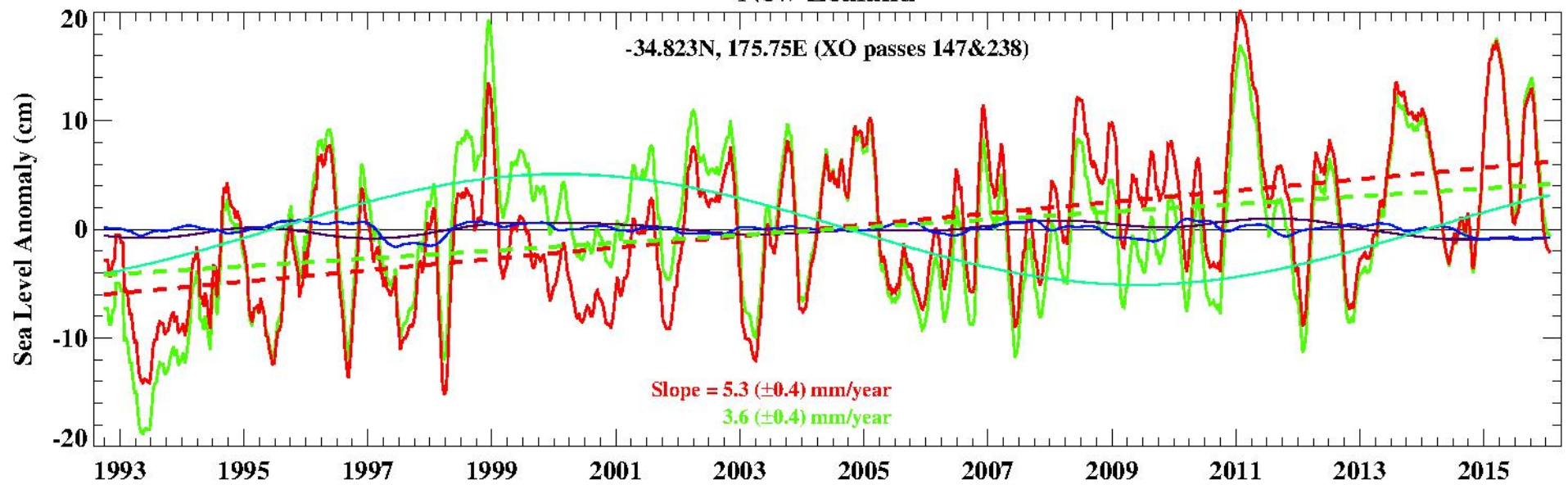


39.216N, 128.98E (XO passes 51&112)

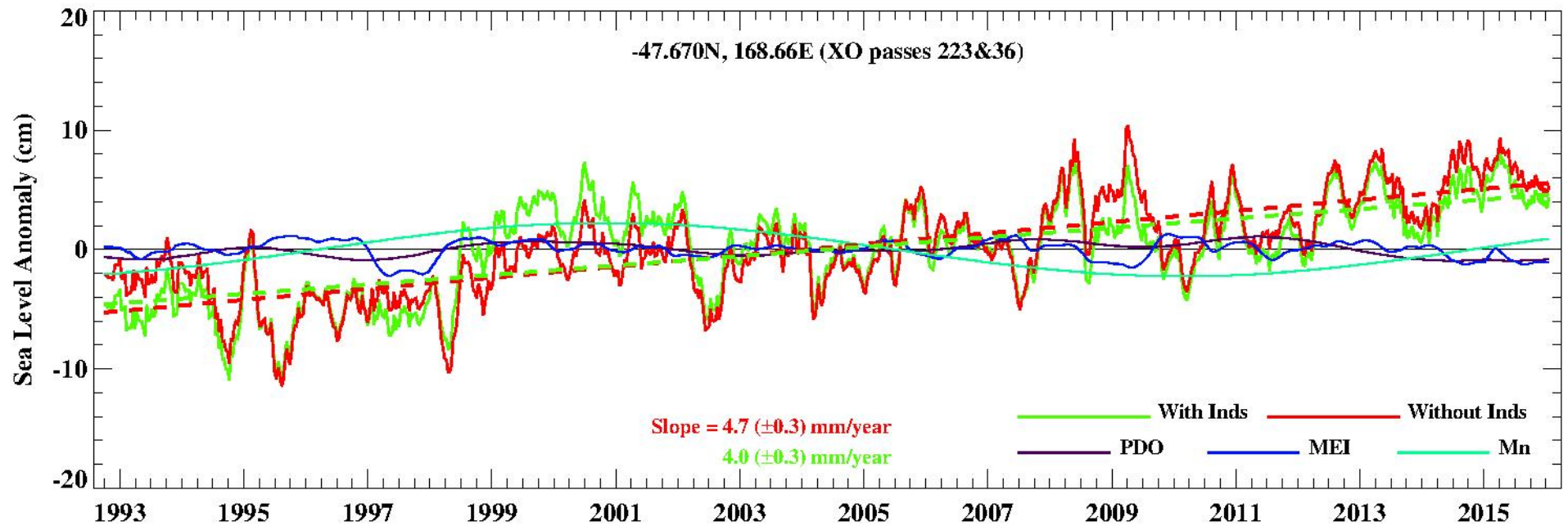


New Zealand

-34.823N, 175.75E (XO passes 147&238)

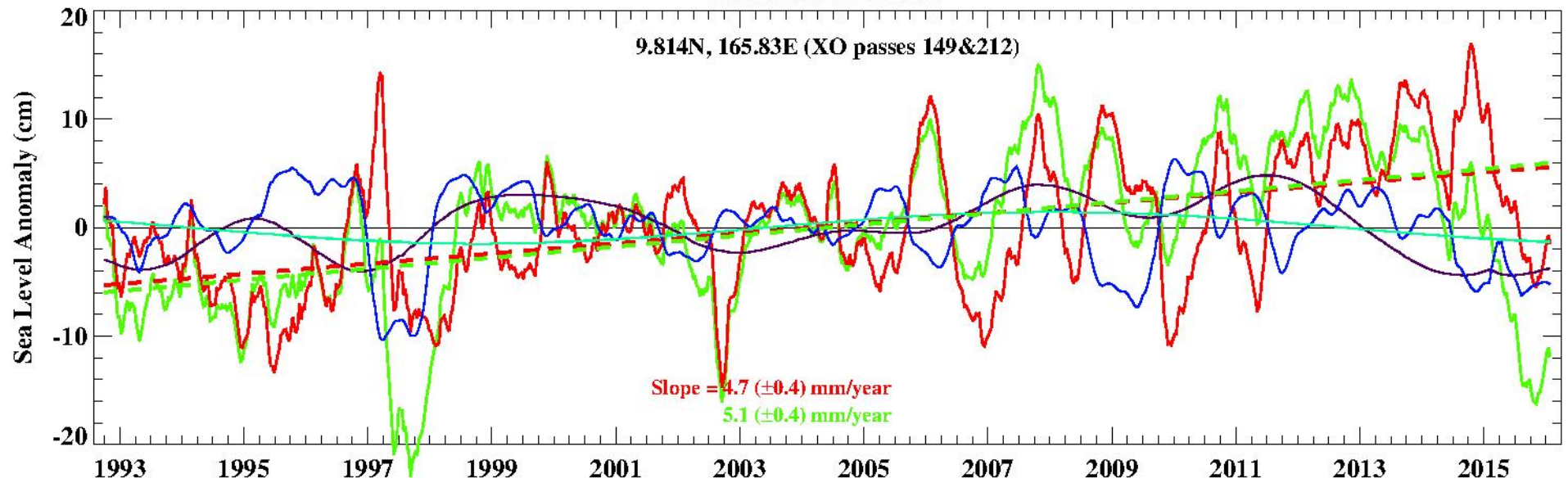


-47.670N, 168.66E (XO passes 223&36)

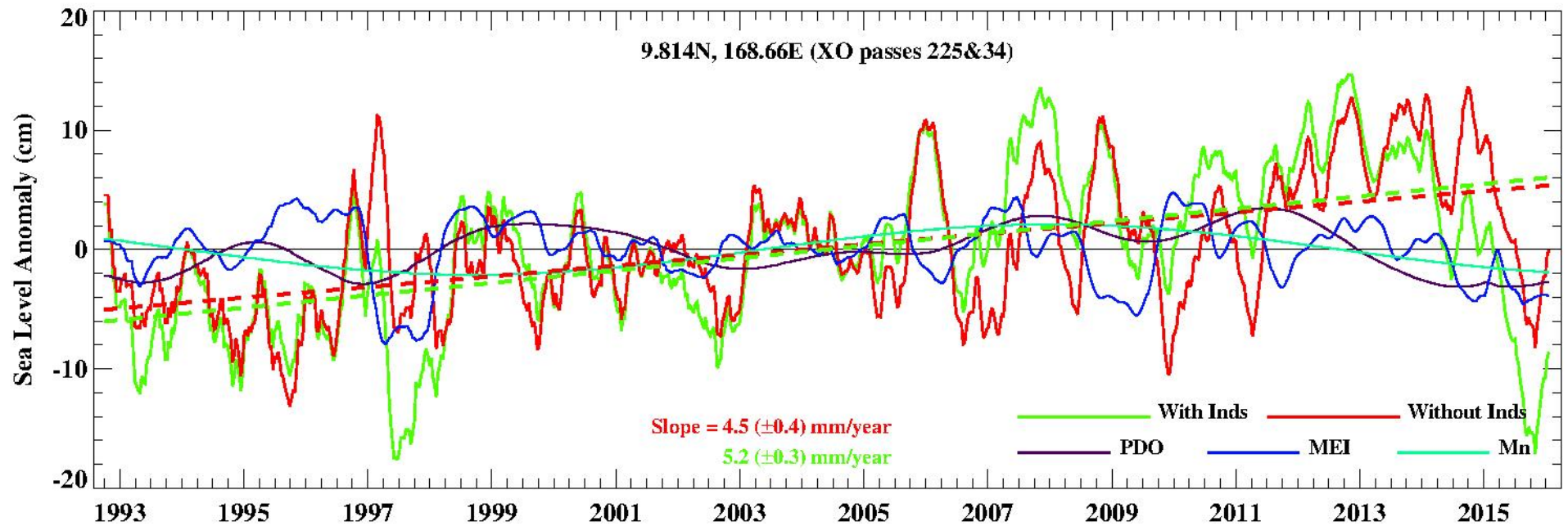


Marshall Islands

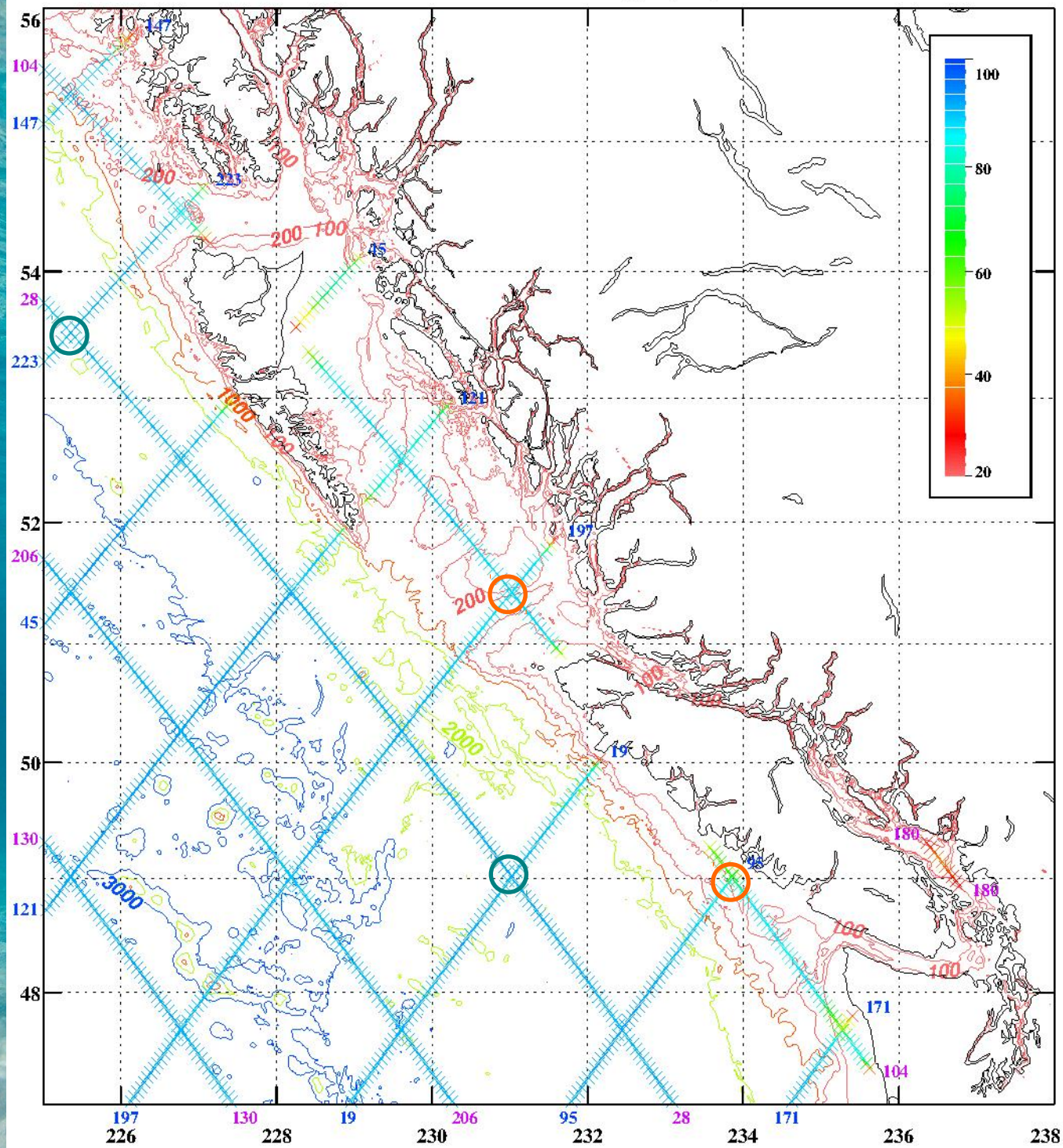
9.814N, 165.83E (XO passes 149&212)



9.814N, 168.66E (XO passes 225&34)

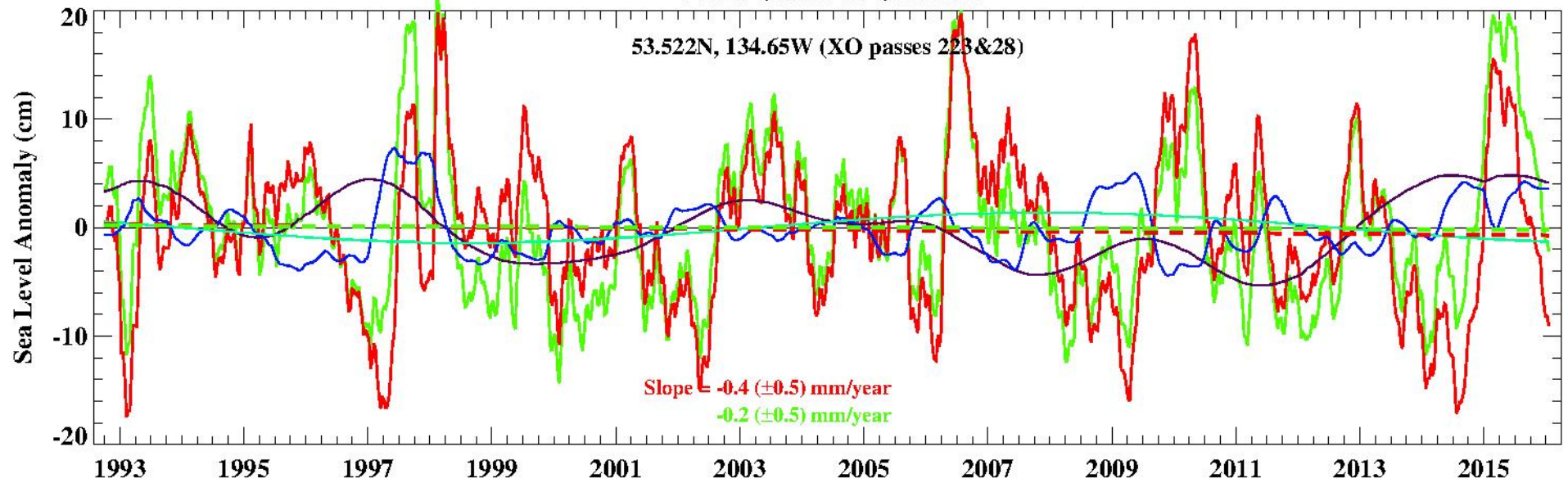


TPJ data coverage (%)

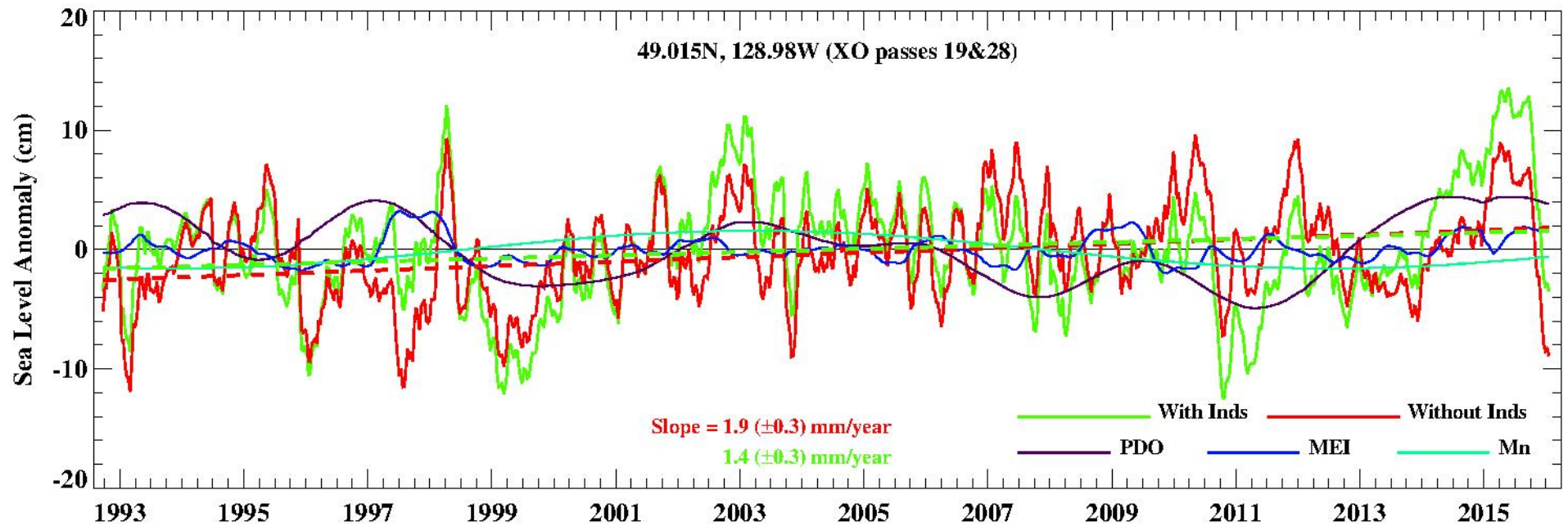


West Coast of Canada

53.522N, 134.65W (XO passes 223&28)

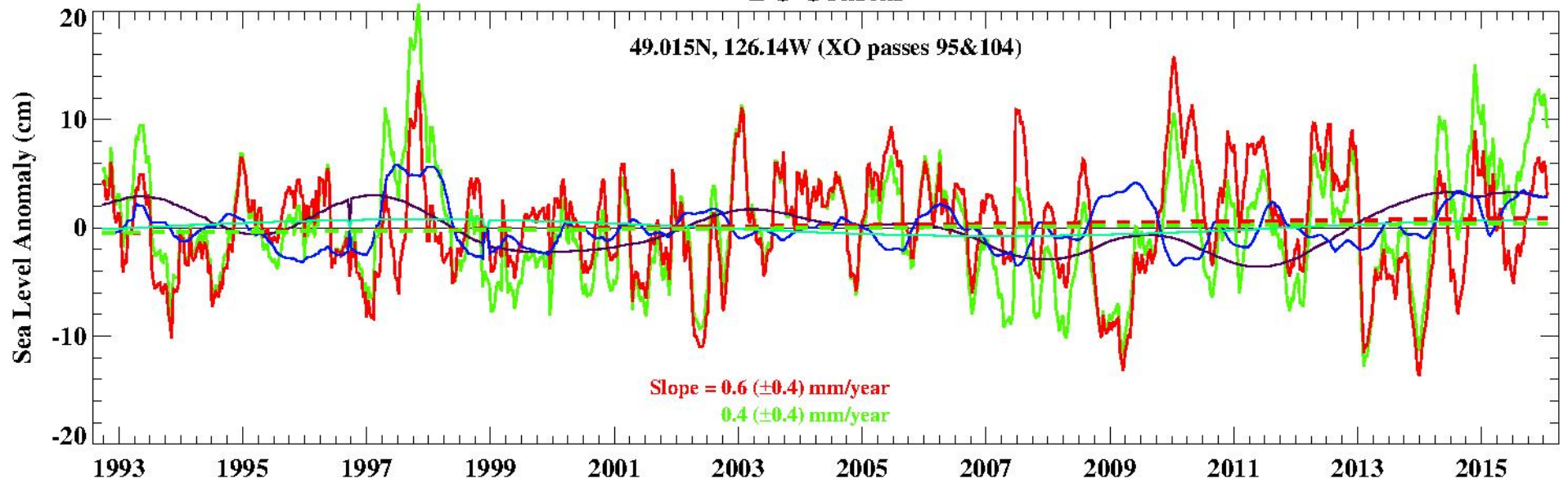


49.015N, 128.98W (XO passes 19&28)

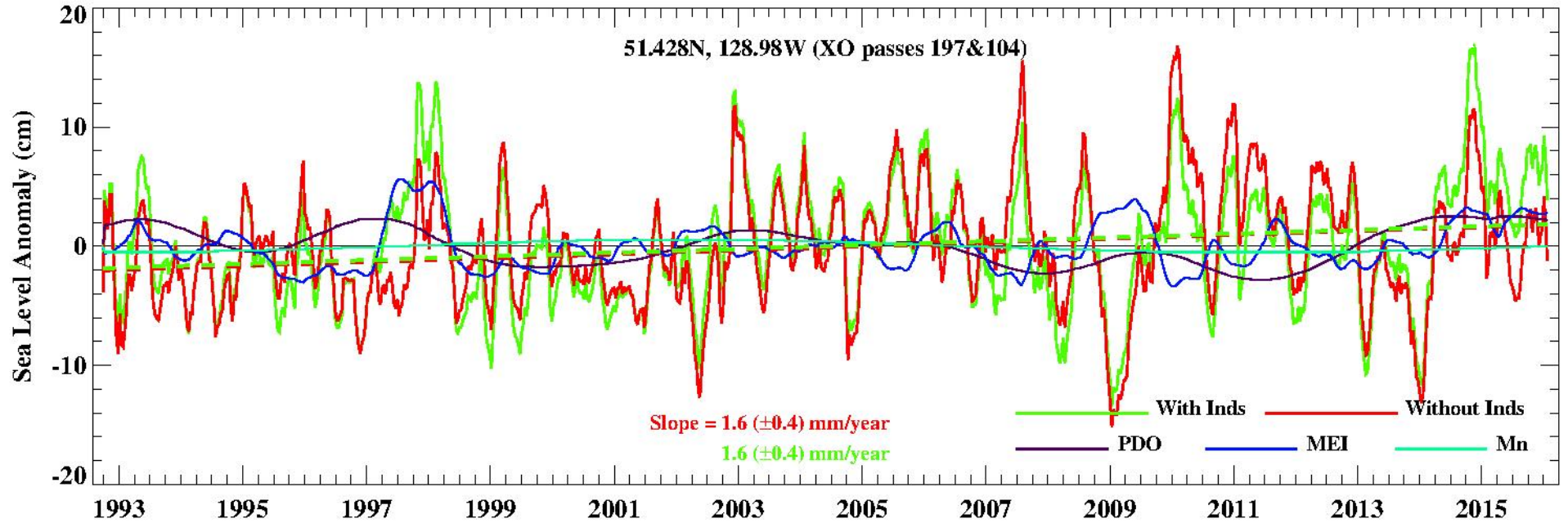


BC Coastal

49.015N, 126.14W (XO passes 95&104)



51.428N, 128.98W (XO passes 197&104)





A diagram illustrating two types of sea level measurements over an underwater scene. The background is a photograph of the ocean surface and seabed. A horizontal dashed orange line represents the sea surface, and another horizontal dashed orange line represents the sea bottom. A vertical yellow arrow points from the top of the image down to the sea surface line, labeled 'Absolute Sea Level (ASL, from altimeters)'. Another vertical yellow arrow points from the sea surface line down to the sea bottom line, labeled 'Relative Sea Level (RSL)'. The text 'Sea surface' is placed just below the first dashed line, and 'Sea bottom' is placed just above the second dashed line.

Absolute Sea Level (ASL, from altimeters)

Sea surface

Relative Sea Level (RSL)

Sea bottom

To get RSL from TPJ data, we must add the relative motion of sea bottom (MSB) which may be due to the Glacial Isostatic Adjustment (GIA), tectonic processes (TP), land subsidence (LS)m etc.

However, we do not know the contributions to MSB from these terms well enough, except for GPS data at the coast. We therefore rely on the models.

The following few slides show our initial attempt to include the Peltier et al. (2015) GIA model results in the TPJ altimeter derived sea level trends.

Peltier et al., J. Geophys. Res. (2015): Space geodesy constrains ice age terminal deglaciation: The global ICE-6G_C (VM5a) model.

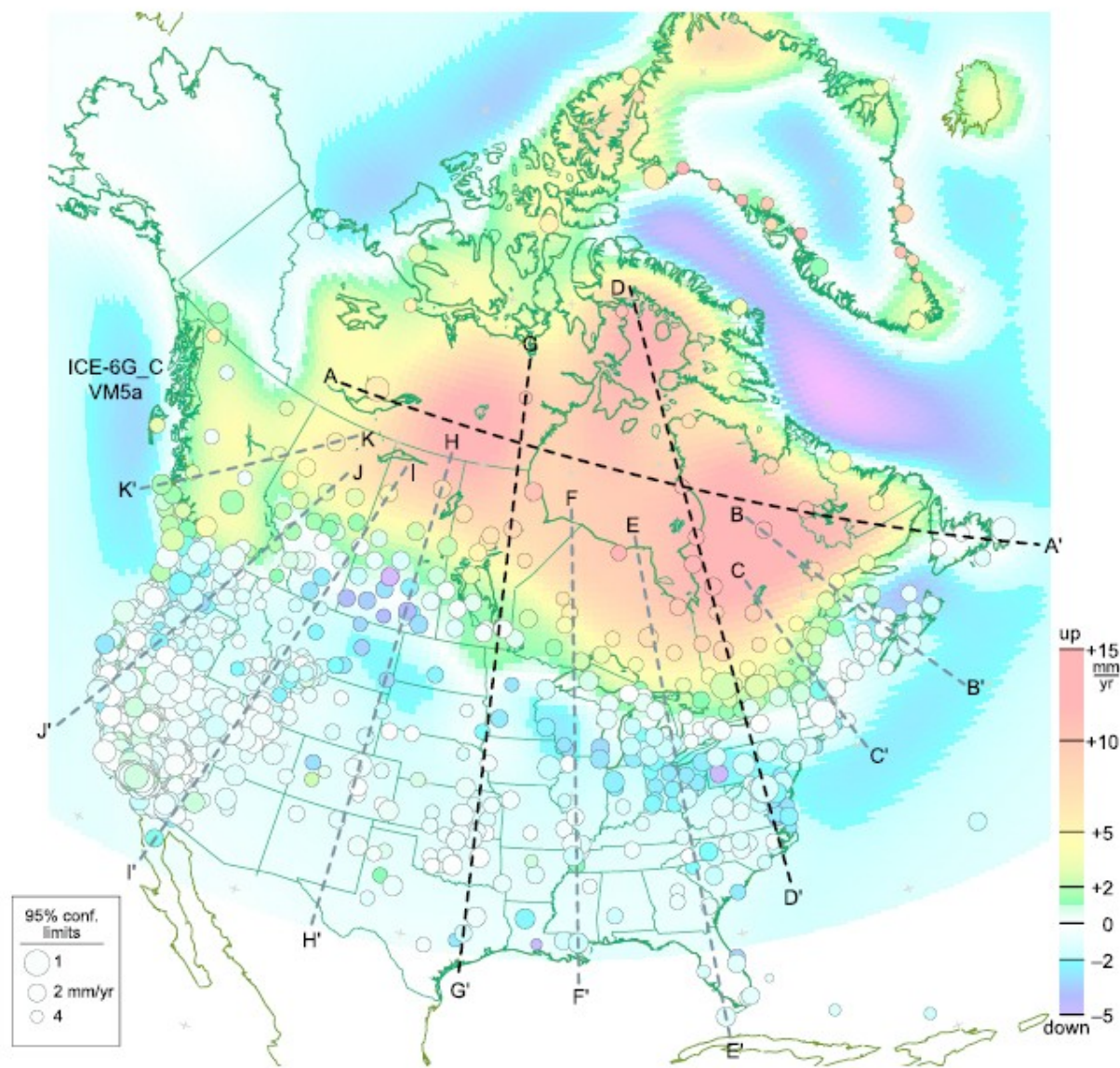


Figure 4. The predicted present-day rate of vertical motion of the crust for the ICE-6G_C (VM5a) model of the global glacial isostatic adjustment process is represented by the background map in which amplitude in mm/yr is represented by the color bar. Superimposed upon this map are the locations of the sites, shown as the open circles, from which GPS measurements of vertical motion are available. The radii of these circles are inversely proportional to the standard error of the individual measurements. Also shown are the traverses along which comparisons are shown in Figures 5a–5c between the predictions of several of the available models including the new model ICE-6G_C (VM5a).

(from Peltier et al., *J. Geophys Res.* 2015)

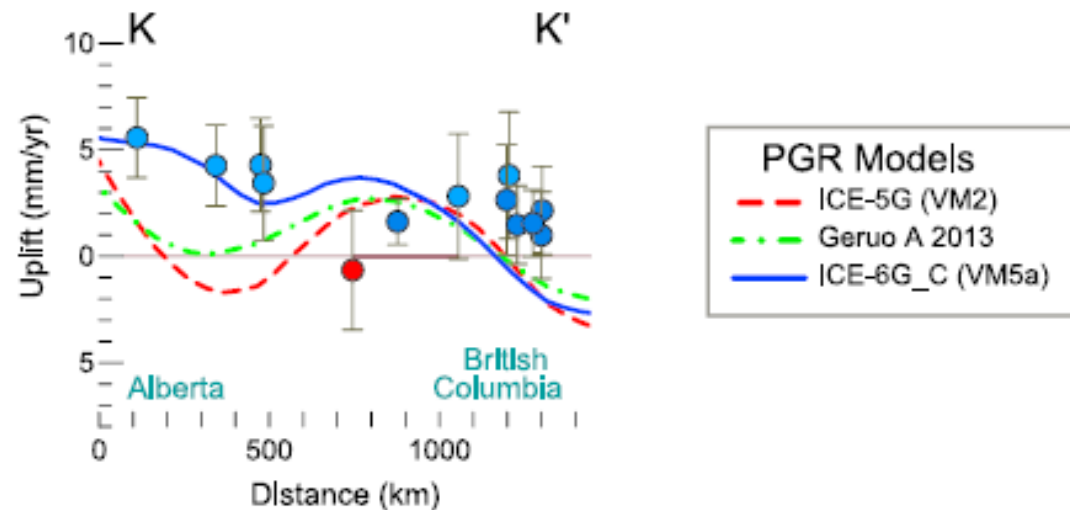
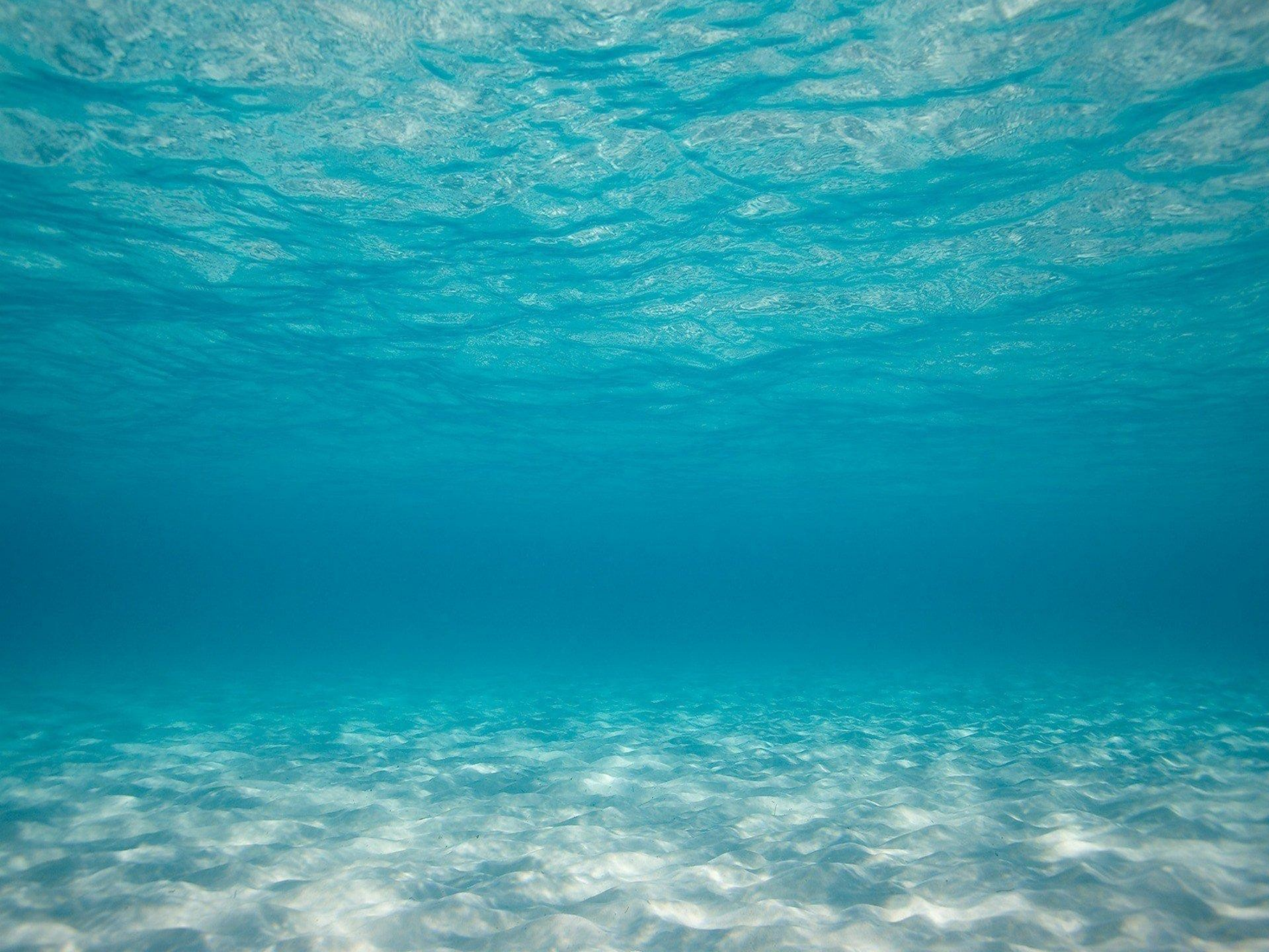


Figure 5. (continued)

Figure 5. (a) Comparisons between GPS observed and GIA model predicted rates of vertical motion along the traverses AA', DD', and GG' shown in Figure 4. Blue sites are from locations that were once ice covered, red sites were never ice covered, and white sites are sufficiently far removed from the line of the traverse that misfits to the data may be ascribable to that source of error. (b) Same as Figure 5a but for the traverses BB', CC', EE', and FF'. (c) Same as Figure 5a but for the traverses HH'–KK'.

(bottom panel of Fig. 5c from Peltier et al., *J. Geophys Res.* 2015)



Conclusions

The 23.4-year long TPJ timeseries are now long enough to compute the lunar nodal tide (Mn) and to estimate its contribution to the analyzed Absolute Sea Level (ASL) trends.

Notwithstanding the cross-correlations, it is also useful to include climate index timeseries (such as PDO and MEI) in the Versatile Analysis of sea level observations.

This is quite evident when comparing sea level trends at TPJ cross-overs with and without the contributions from climate indices.

Inclusion of GIA increases the analyzed global mean sea level trend from 2.8 to 3.1 mm/year. However, on smaller regional scales our knowledge of sea bottom motion is not adequate and may lead to erroneous trends.

This further underscores the utmost importance of the coastal tide gauges, which together with ASL from satellite altimeters and data from nearby GPS stations, allow to analyze the Relative Sea Level (RSL) changes at the coast.